

Use of precipitation radar in hydrologic modelling in Norway

Perspectives for a new research project

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Outline

- New project
- Previous projects
- Analysis of Norwegian radar data
- Bayesian combination of radar and gauge data
- Gaussian random fields for conditional simulations

New project:

- Operational use of weather radar, for EBL 2009-2011
- Cooperation with met.no, Cemagref and NTNU
- Continuing development of the radar-gauge assimilation, hopefully merging with the simulation approach
- Analysing reflectivity data from the radar for assimilation into fine scale meteorological models.
- Error models for radar data, and the value of gauges in improving radar data quality
 - Hourly data and temporal covariance structure
 - Special focus on mountainous areas
 - Effect on precipitation validation on reflectivity data quality

Recent projects

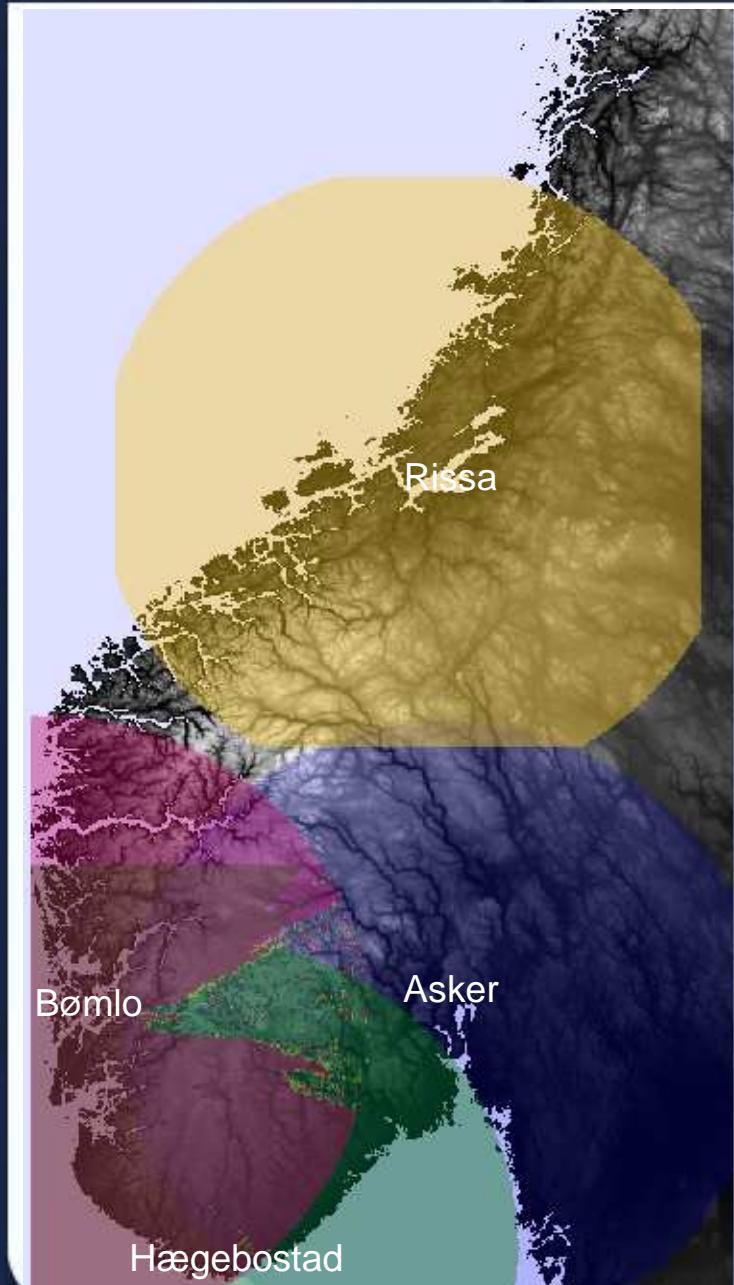
- Developed and implemented conditional simulation method as an alternative to Inverse distance
- Development of a distributed model framework
- Improved simulation of snow processes
- Implemented a Kriging approach, focusing on station selection and variable uncertainty
- Started evaluating weather radar maps, and implementing Todini's (2001) routine for Bayesian assimilation of radar and gauge data
- Also focusing on the use of gridded meteorological forecasts

Ongoing and planned projects

- "The Hydrological Crystal Ball" for Statkraft 2008-2011
 - A full range project on uncertainty in operational forecasts
 - Target: Decision supporting information for power traders, includes estimating the economic value of information
 - Error models for measured data:
 - Precipitation and temperature
 - Discharge and naturalised flow series.
 - Simulation using meteorological ensemble forecasts
 - Parameter uncertainty estimation
 - Error models for the HBV lumped model and a distributed model
 - One of several projects on implementing distributed models

Precipitation data

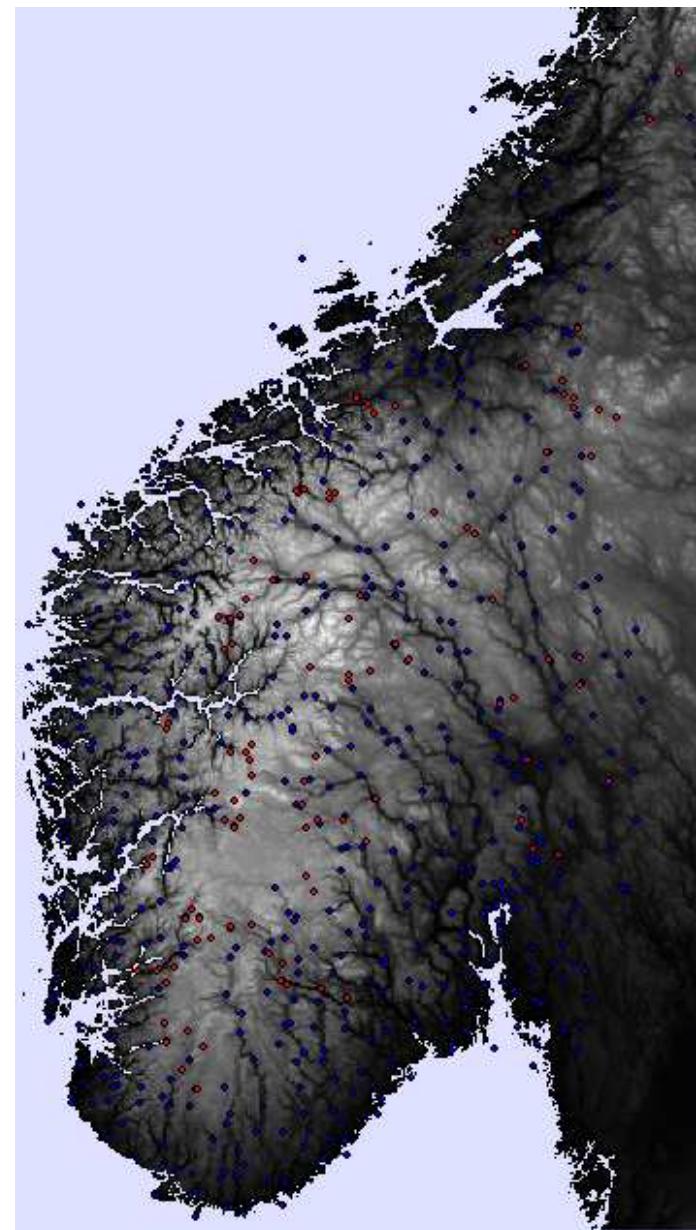
| | Radar | Gauge |
|---------------------|--------------------|----------------------|
| Spatial support | ~1 km ² | ~100 cm ² |
| Spatial resolution | ~1 km | ~10 km - ~100 km |
| Spatial coverage | Southern Norway | All Norway |
| Temporal support | ~4x1 min? | ~1 h; ~24h |
| Temporal resolution | ~1h | ~1 h; ~24h |
| Temporal coverage | 2000 -> | 1900 -> |
| Bias | Large ? | Wind dependent |
| Uncertainty | Large ? | Small |
| | | |



Radar data:

- Rissa (Jan 2006- April 2008)
- Bømlo (April 2007-April 2008)
- Asker (April 2007-April 2008)
- Hægebostad (April 2007-April 2008)

Daily data accumulated from hourly data



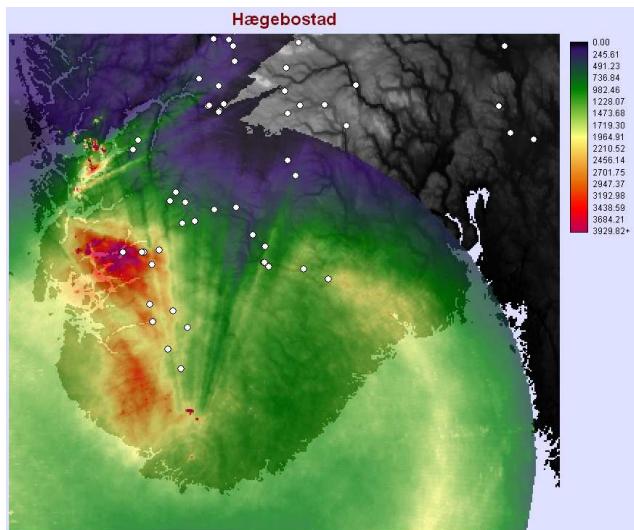
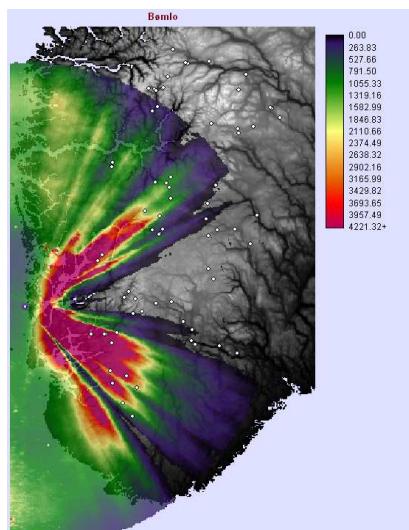
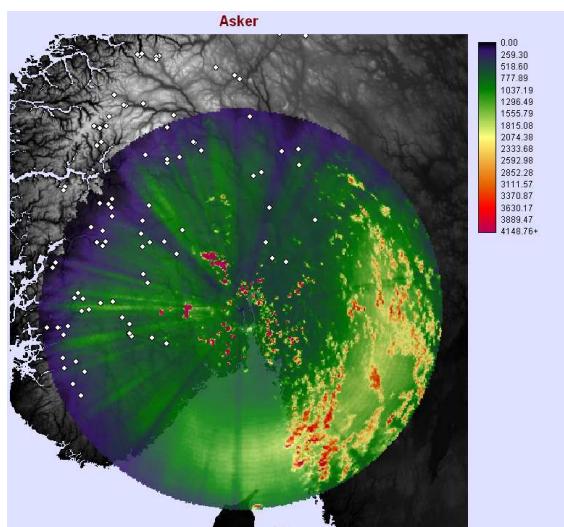
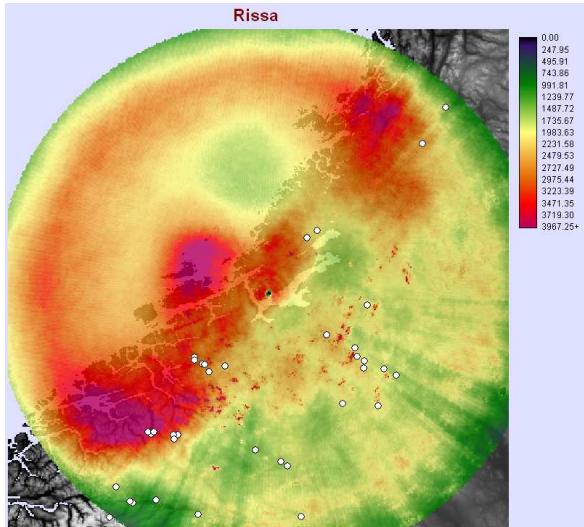
Precipitation gauge data for evaluation:

- **dependent stations**
 - Norwegian met. office
 - already used in adjusting the radar precipitation
- **independent stations**
 - hydro power companies

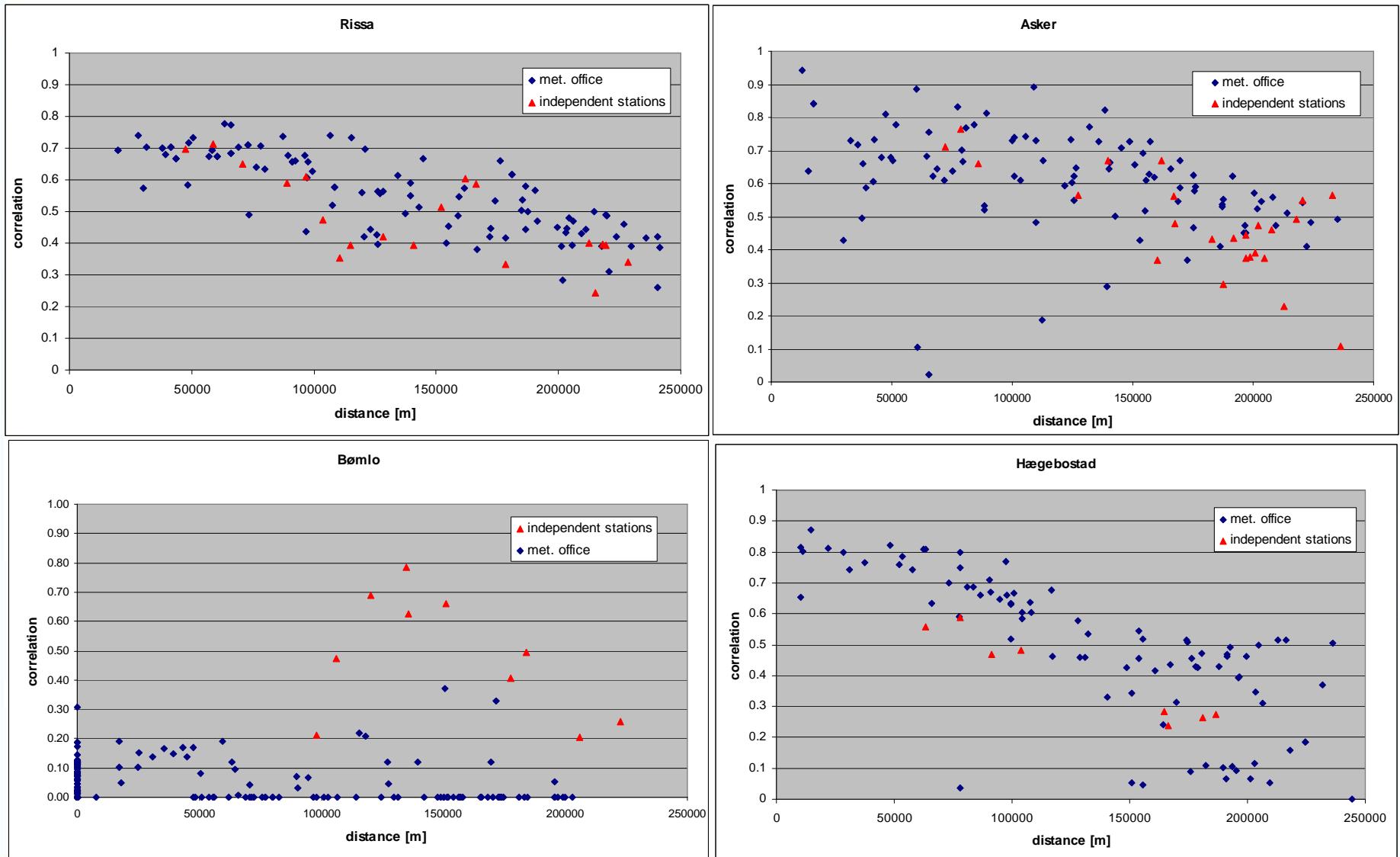
Evaluation against gauge stations

- correlation
- average values
- Nash-Sutcliffes efficiency criteria R^2

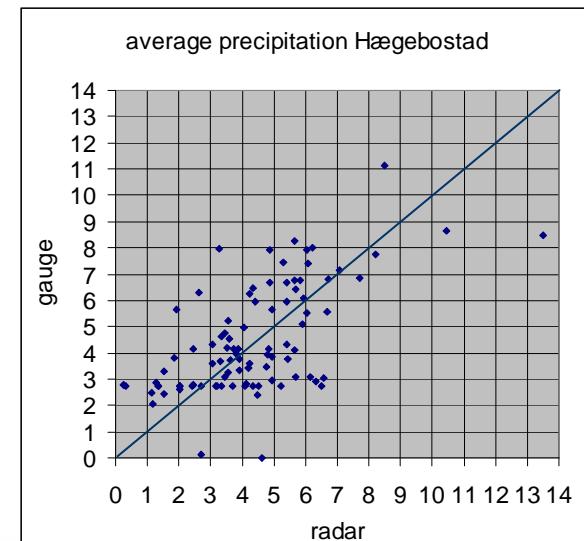
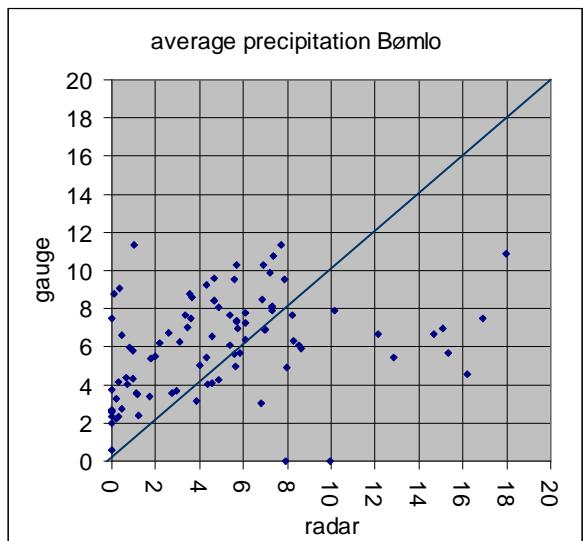
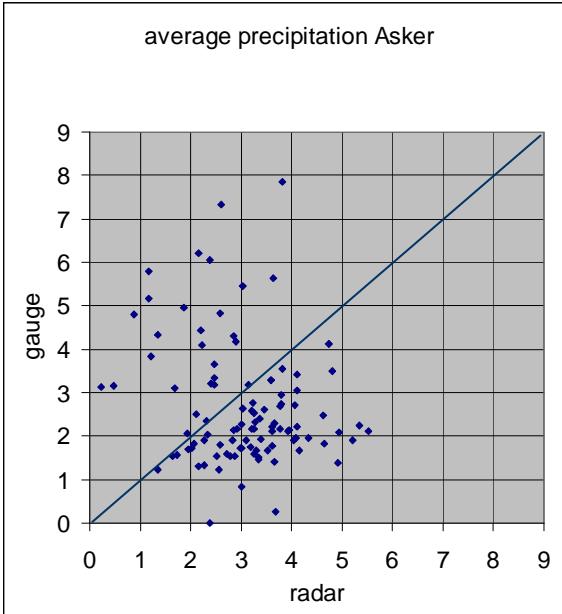
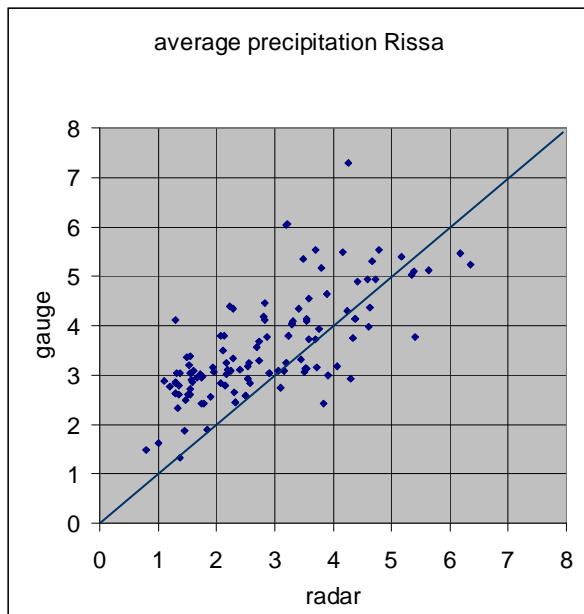
Sum radar precipitation



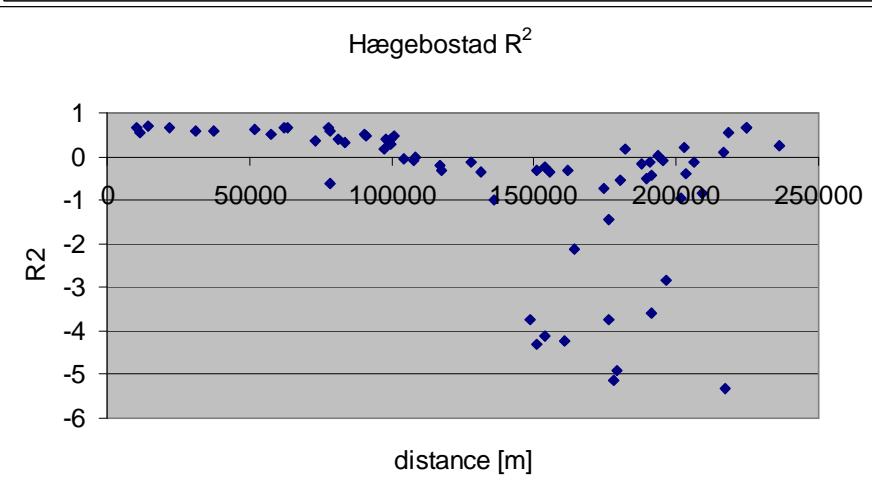
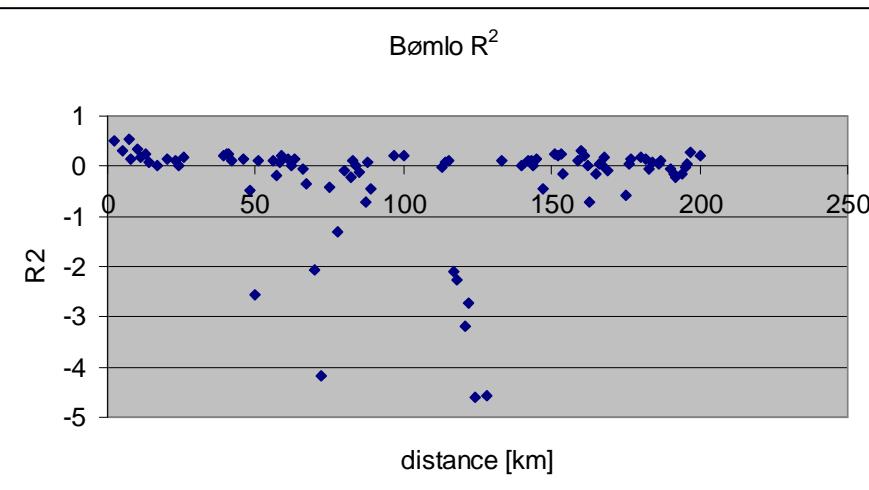
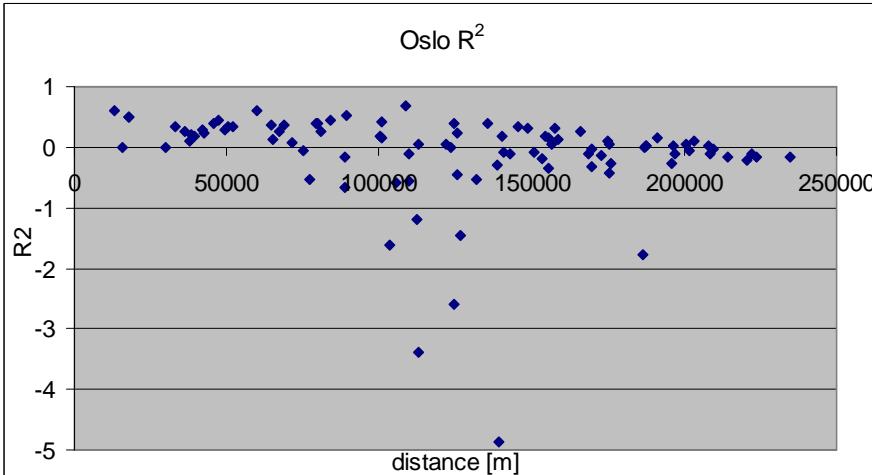
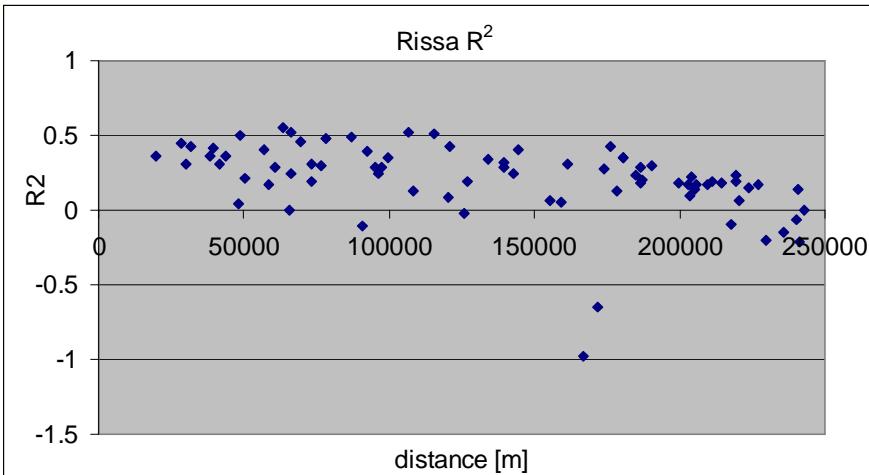
Correlation



Average values



Nash-Sutcliffe R^2



Bayesian Combination method

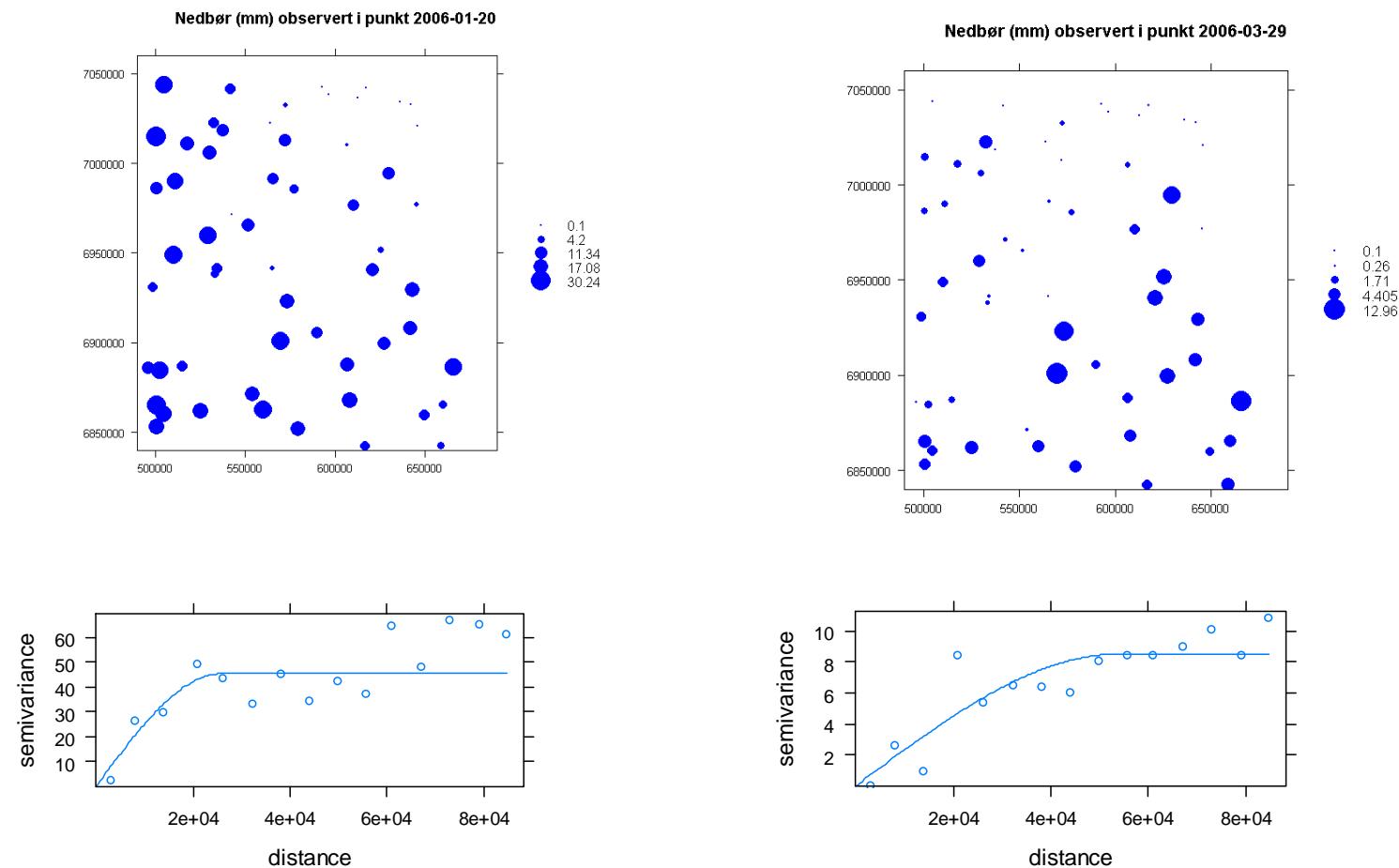
- Interpolation of point precipitation to grid-squares 1x1 km²
- Estimate the bias and variance of the radar precipitation
- Combine the interpolated grid and the radar based on Kalmar gain.

Interpolation to grid squares

- Using block Kriging
- Estimate point semivariogram
- Estimate covariances between points and grid-squares
- Expectation and covariance

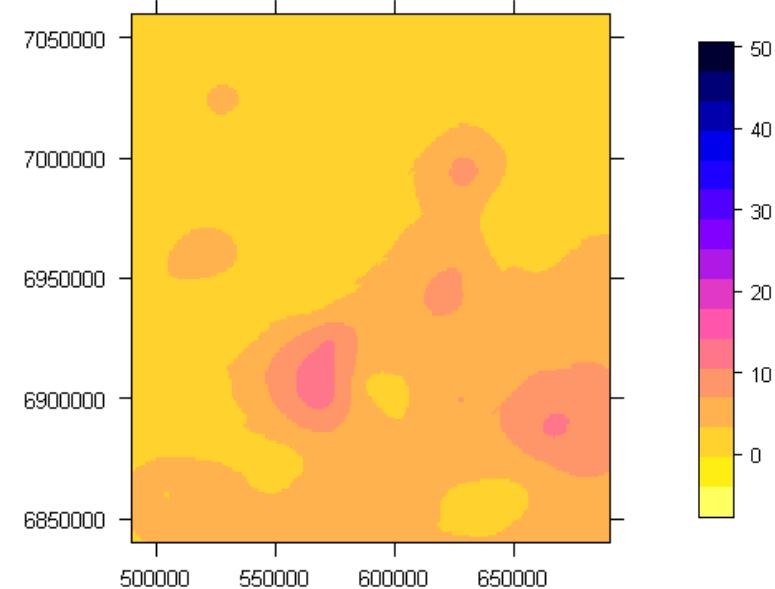
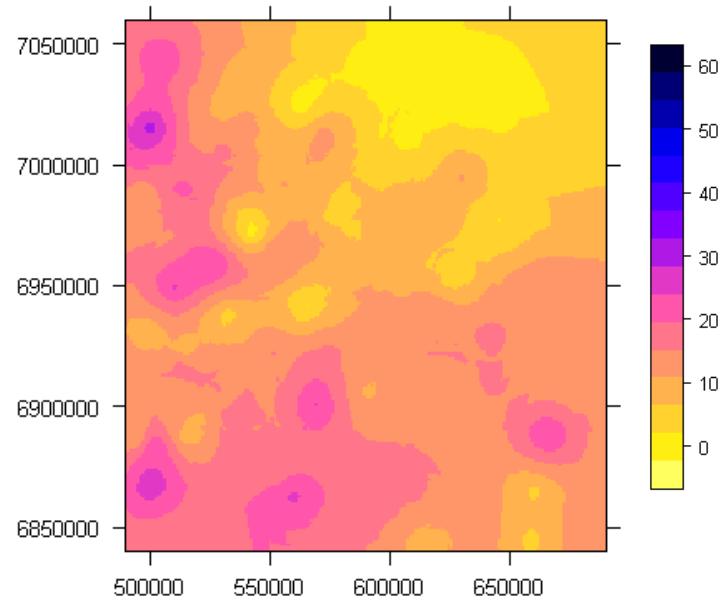
Interpolation to grid squares

20. January 2006 and 29. March 2006



Interpolation to grid squares

20. January 2006 and 29. March 2006

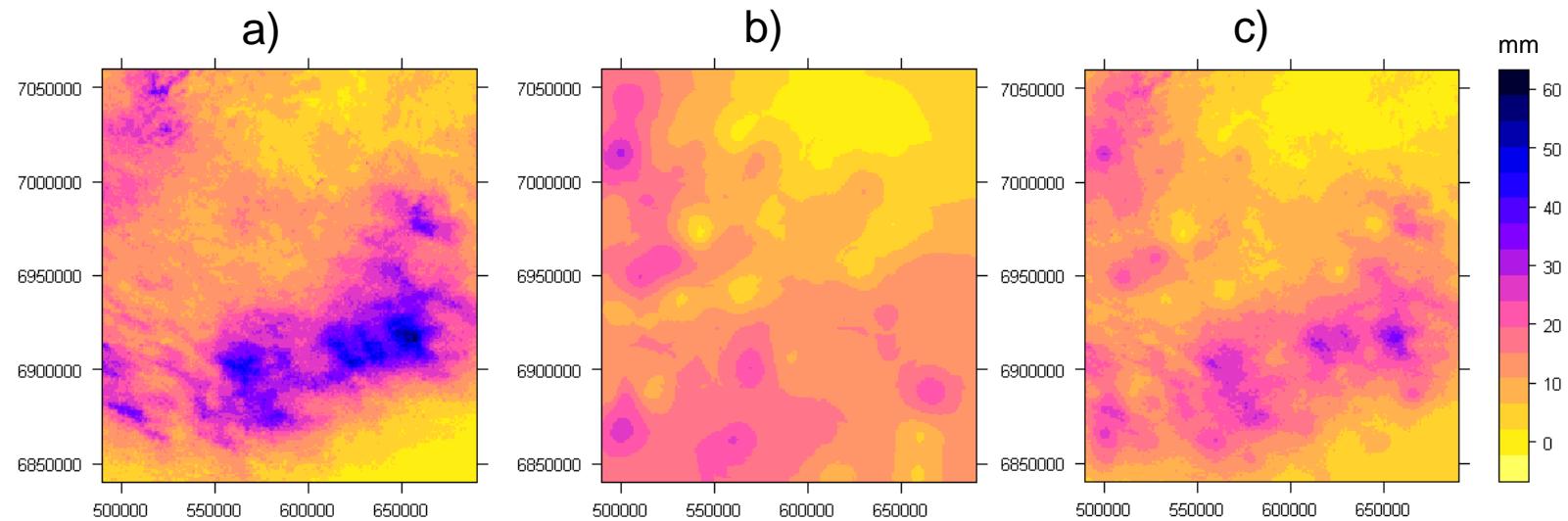


Kalman filter

- Assuming block-kriged values are un-biased and the kriging covariance is known.
- Bias of radar precipitation is average difference between interpolated and radar values
- Exponential semiovriogram is fitted to the differences and the covariance matrix of radar precipitation is estimated from the fitted semiovriogram
- The two data sources are weighted according to their variance.
 - I.e. decreasing differences between the radar and the interpolation, leads to increasing weight to the radar image

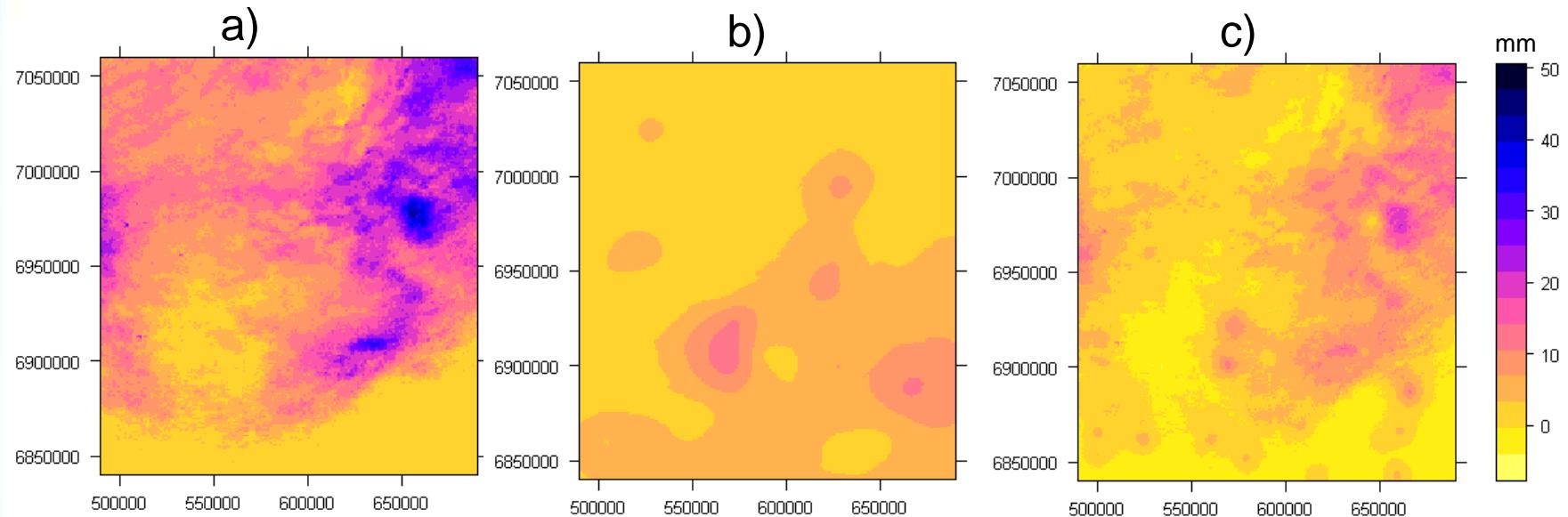
Results

■ 20 january 2006



Results

■ 29 mars 2006



Summary

- Makes a reasonable combination
- Many challenges
 - Zero precipitation.
 - Interpolation
 - Radar
 - Interpolated precipitation will be smooth and not a realistic precipitation-field whereas radar images are more patchy
 - Independent estimates of bias and variance in radar precipitation.
 - Time dependence (advection).
 - Alternative approaches:
 - Use conditional simulations in stead of block-kriging
 - Use radar images in conditional simulations.

Conditional simulation of precipitation

- Provides several possible realisations of precipitation.
- Might obtain both expectation and uncertainty of both catchment precipitation and the distributed precipitation field in a consistent way.
- Accounts for zero precipitation.
- Observed values are reproduced.

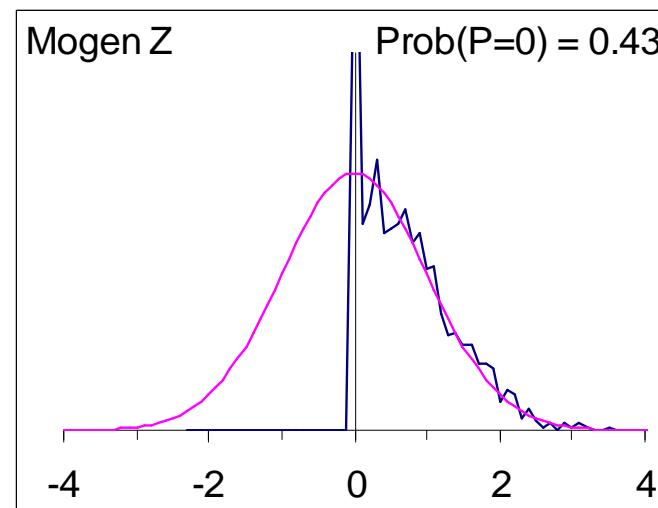
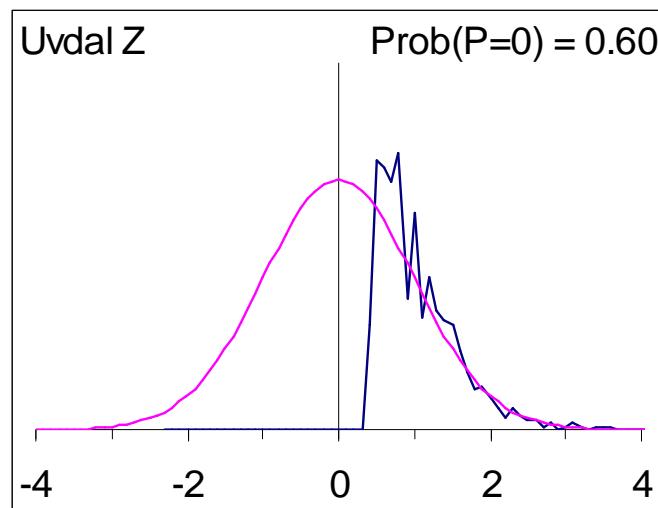
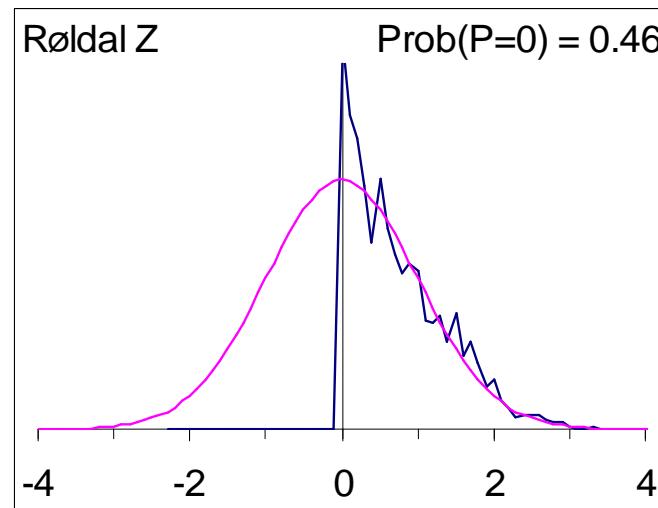
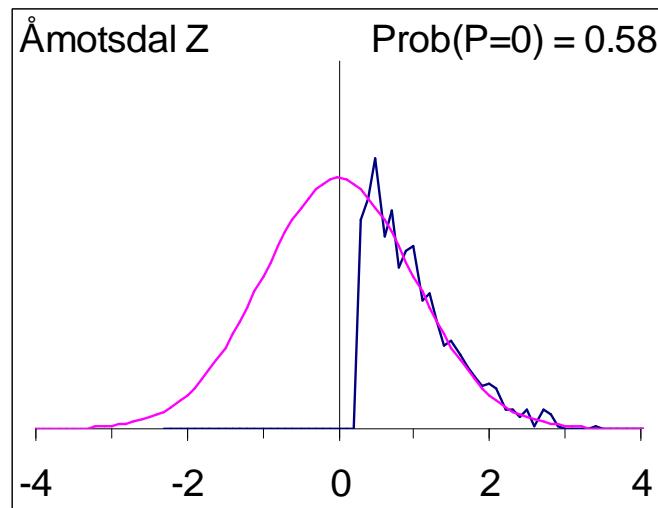
Operational motivation

- The tradition: ~200 lumped HBV models, in each catchment calibrating or subjectively assessing
 - 2 catch deficit compensating parameters
 - 1 elevation gradient
 - 2-3 gauge weights
- In a distributed model needing gridded input data, a less calibration-dominated approach is required.
- Input uncertainty estimation is desirable both for operational runoff prediction and to reduce over-conditioning in calibration

Simulation approach

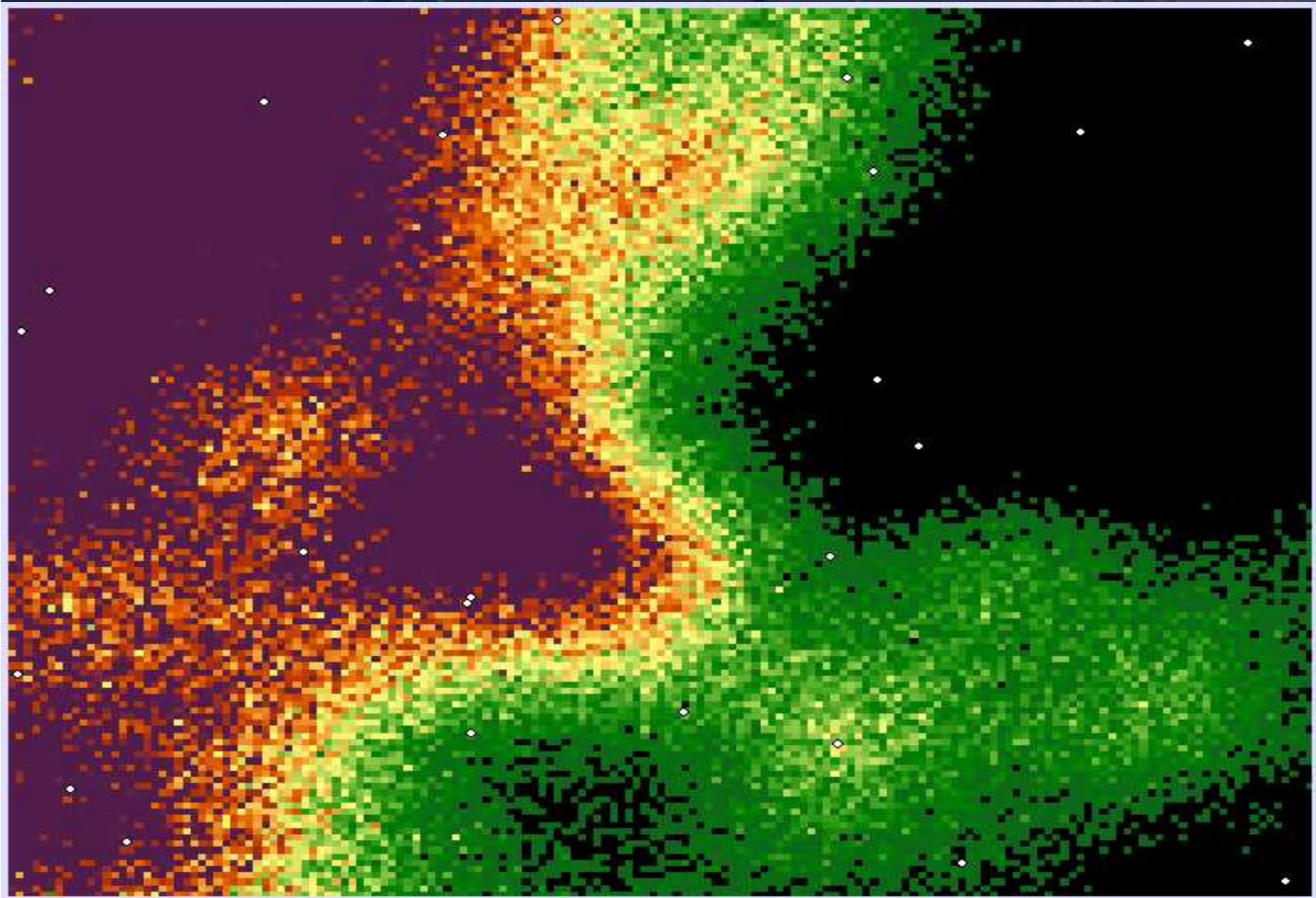
- Preprocessing:
 - For each station positive precipitation is fitted to a gamma distribution and probability of no precipitation, then Interpolate parameters.
- For each day
 - Transform all positive P observations to truncated normally distributed values above the truncation level
 - Initiate $Z < Z_0$ at $P=0$ sites.
 - Conditionally simulate Z at stations observing $P=0$, drawing from a truncated Normal distribution ($Z < Z_0$), given neighbour Z values. This step employs a Markov sequence since also the other 0-observing stations are included in the conditioning set. Repeat n_1 times.
 - Simulate the complete field, conditionally on both the transformed Z values and the current simulated Z values. Repeat step 1 and 2 n_2 times.
 - Back-transform each of the n_2 realisations, using the maps of Z_0 and transformation parameters. Aggregate the desired statistics for the back-transformed P values, for instance $E[P]$, $HPD[P]$, $Var[P]$, $Prob(P>0)$,

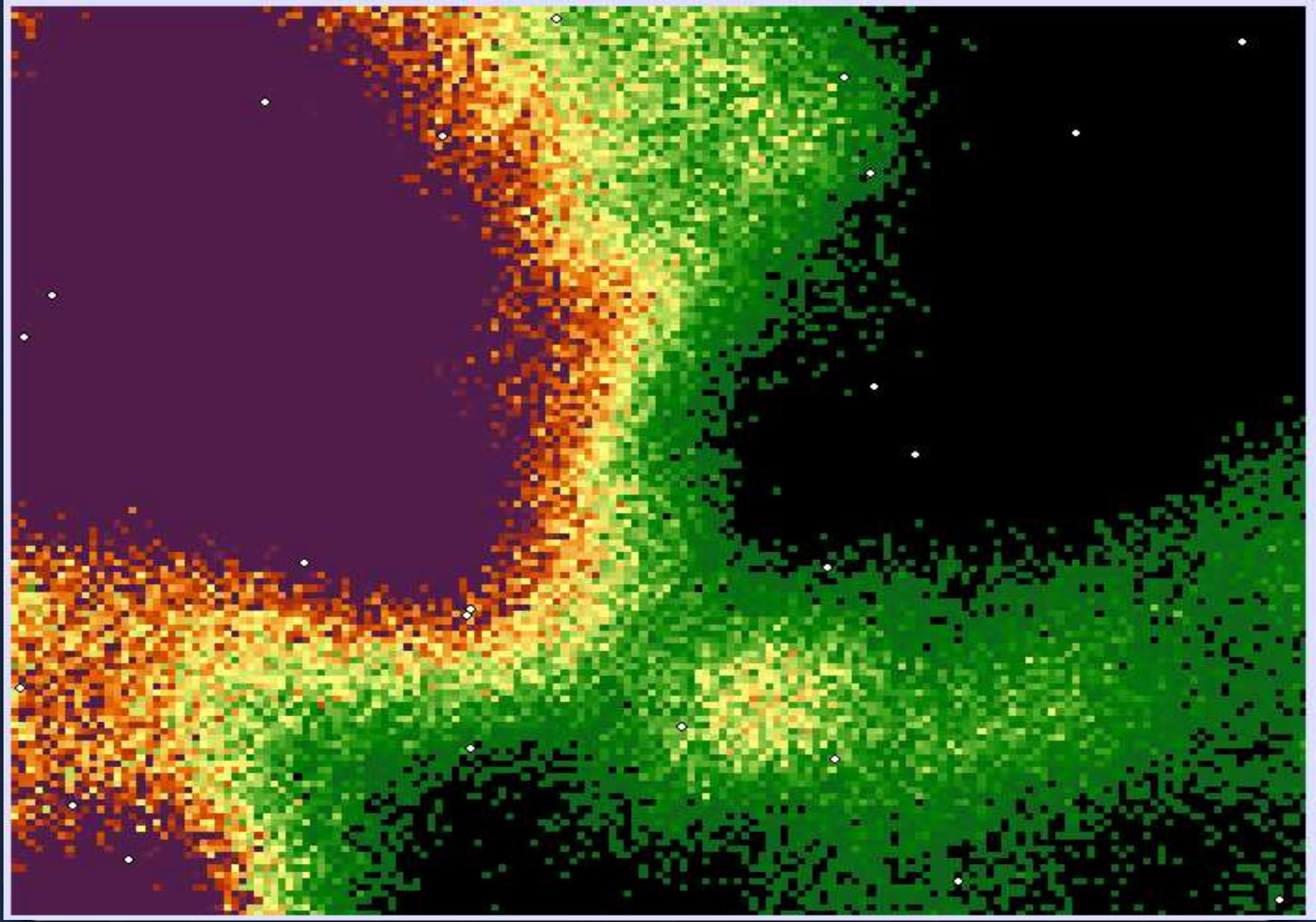
Z standard normal with different truncation level

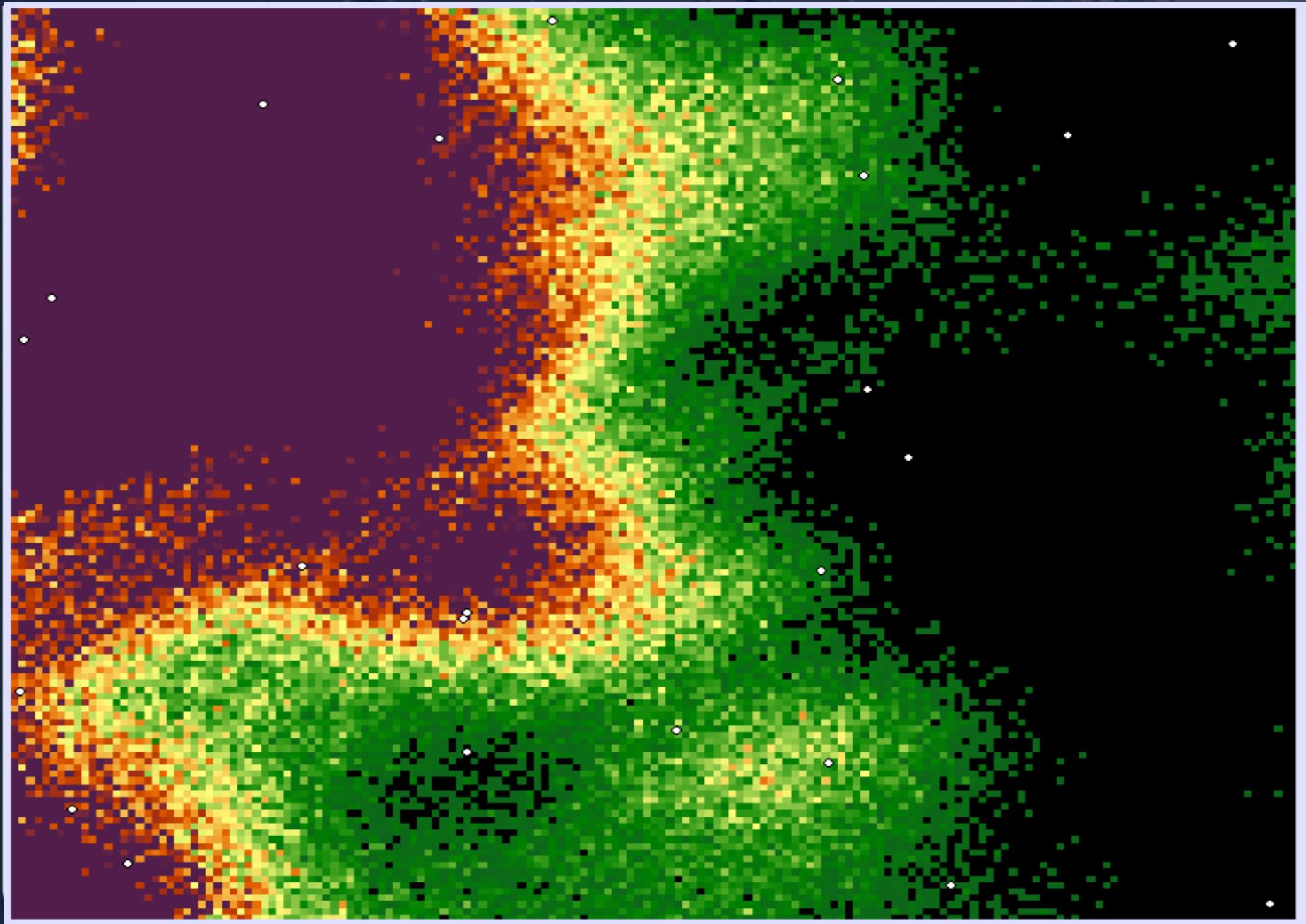


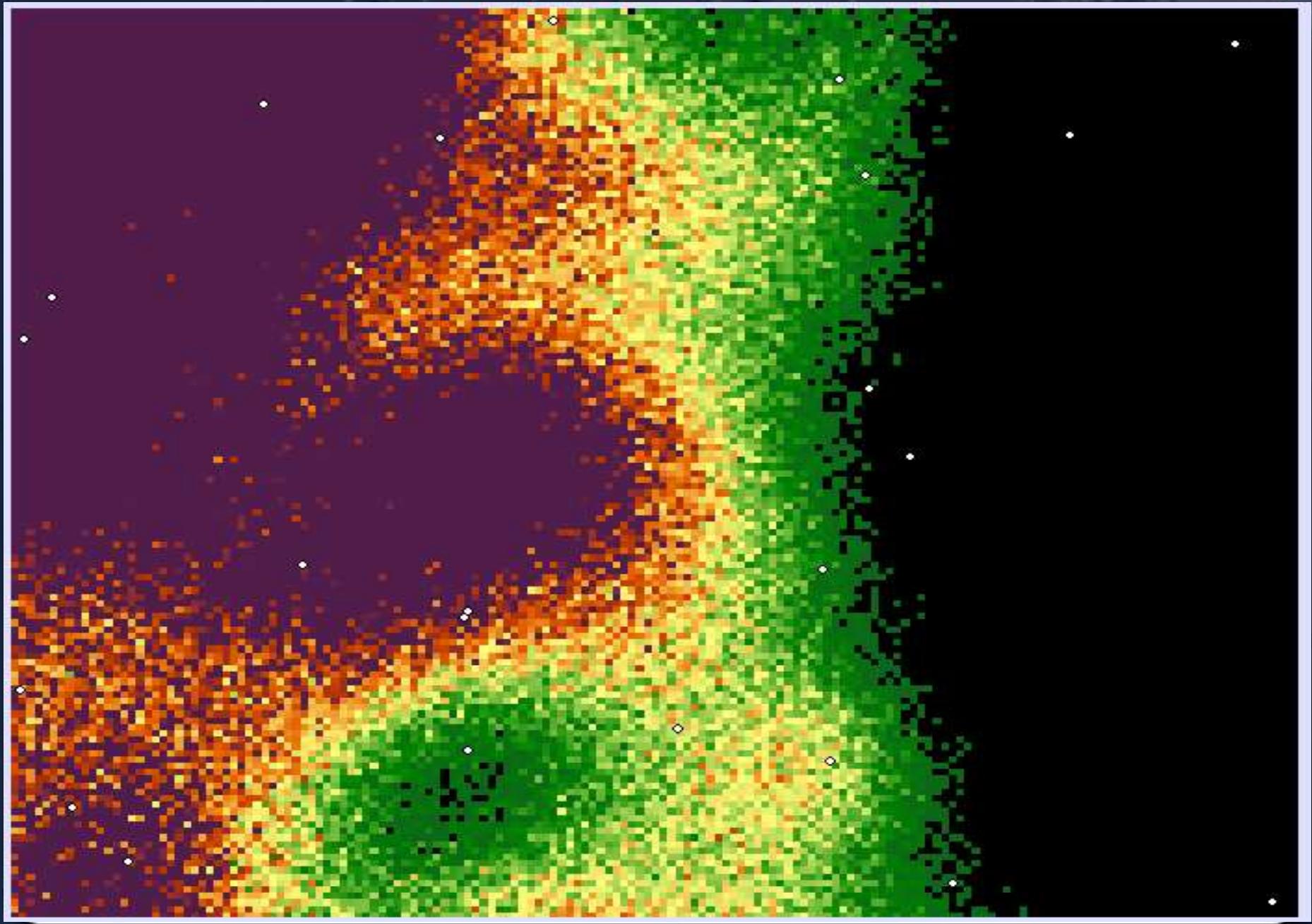
Pragmatic decisions in current version

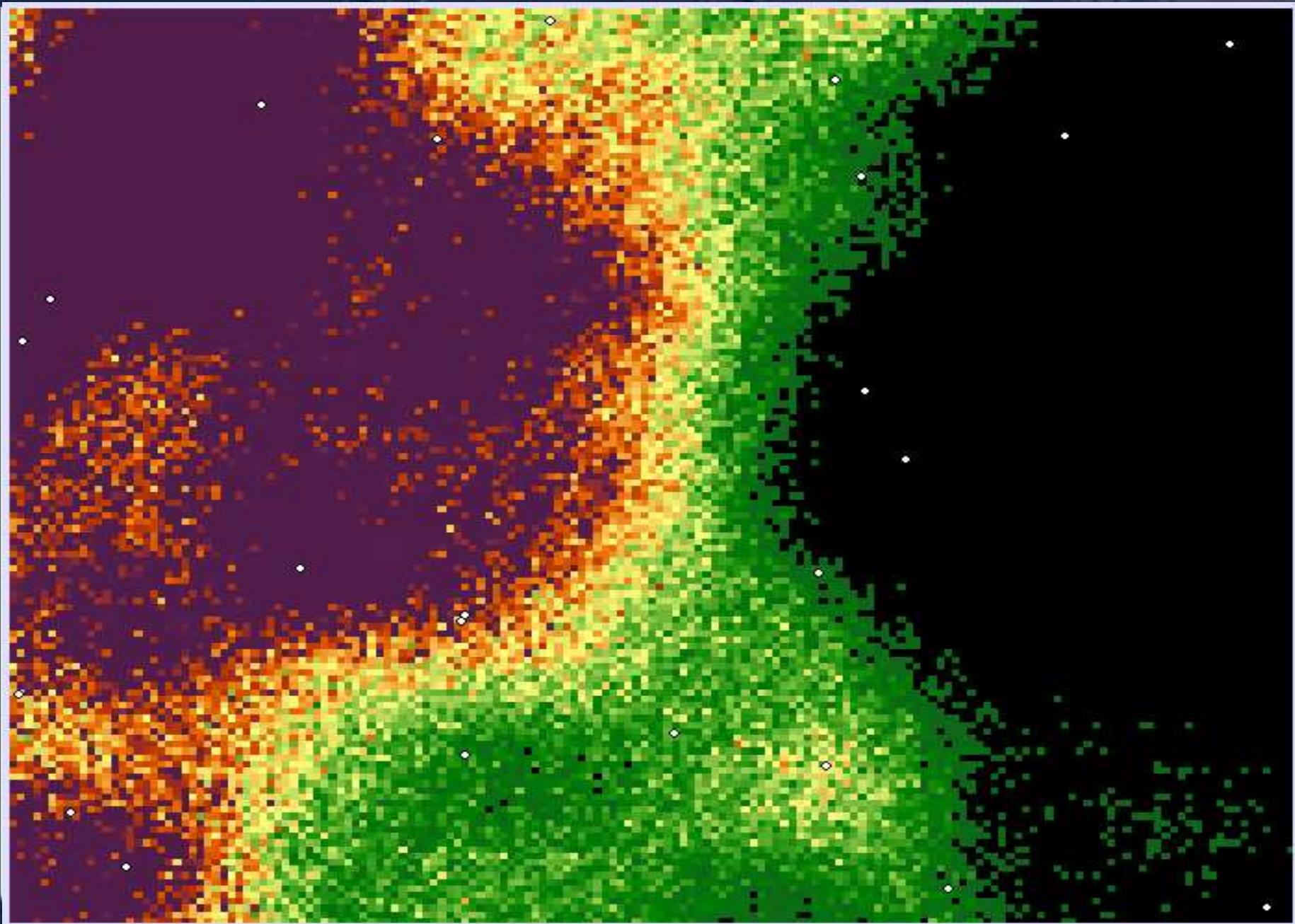
- Estimation of the covariance structure (semivariogram) is external to the model, and can be treated separately.
- A Matern semivariogram model, and a Gamma density model for positive precipitation is assumed
- So far no seasonal or situational conditioning of the assumptions, neither daily estimation of semivariogram (apart from variance scaling).

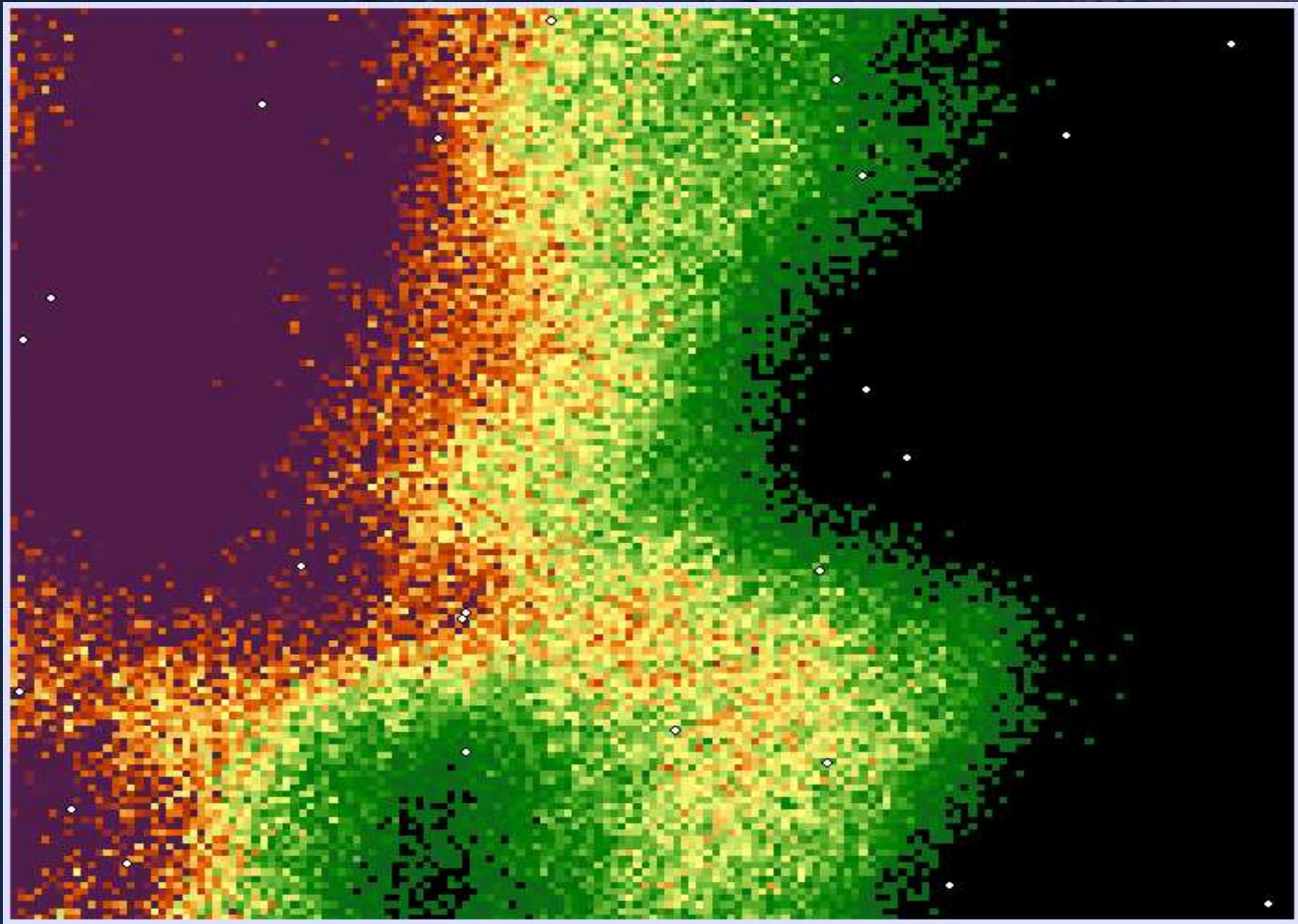


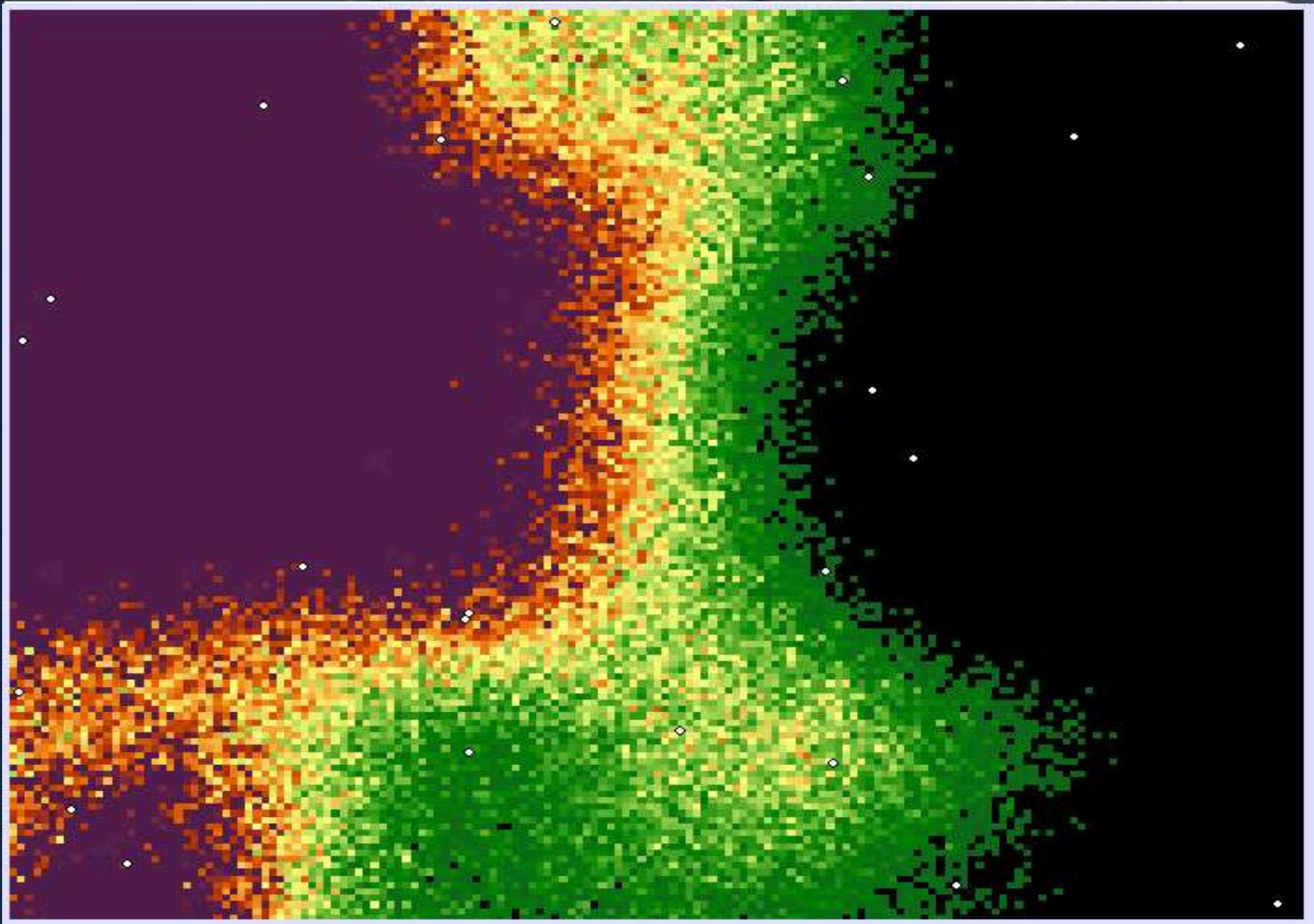


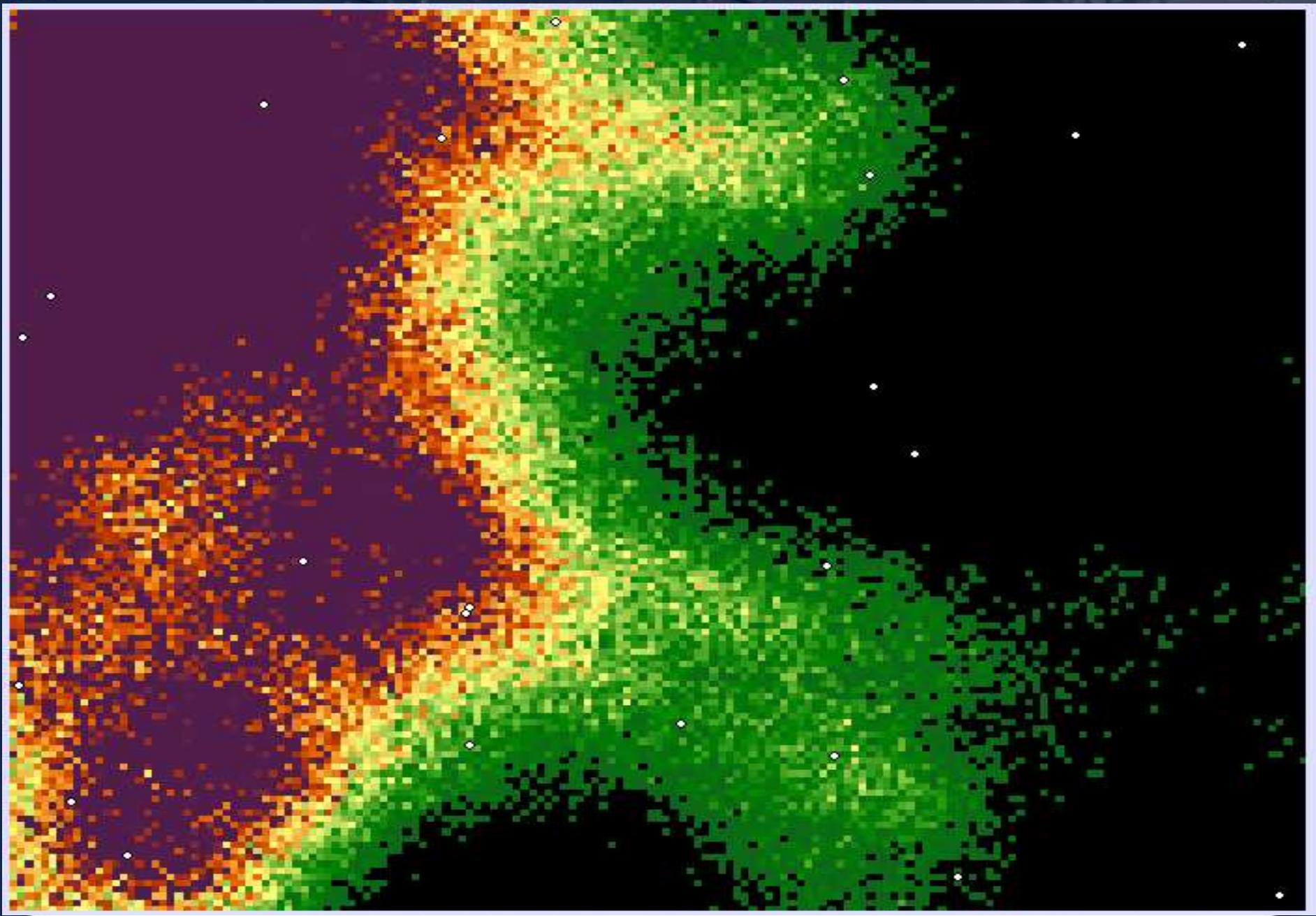


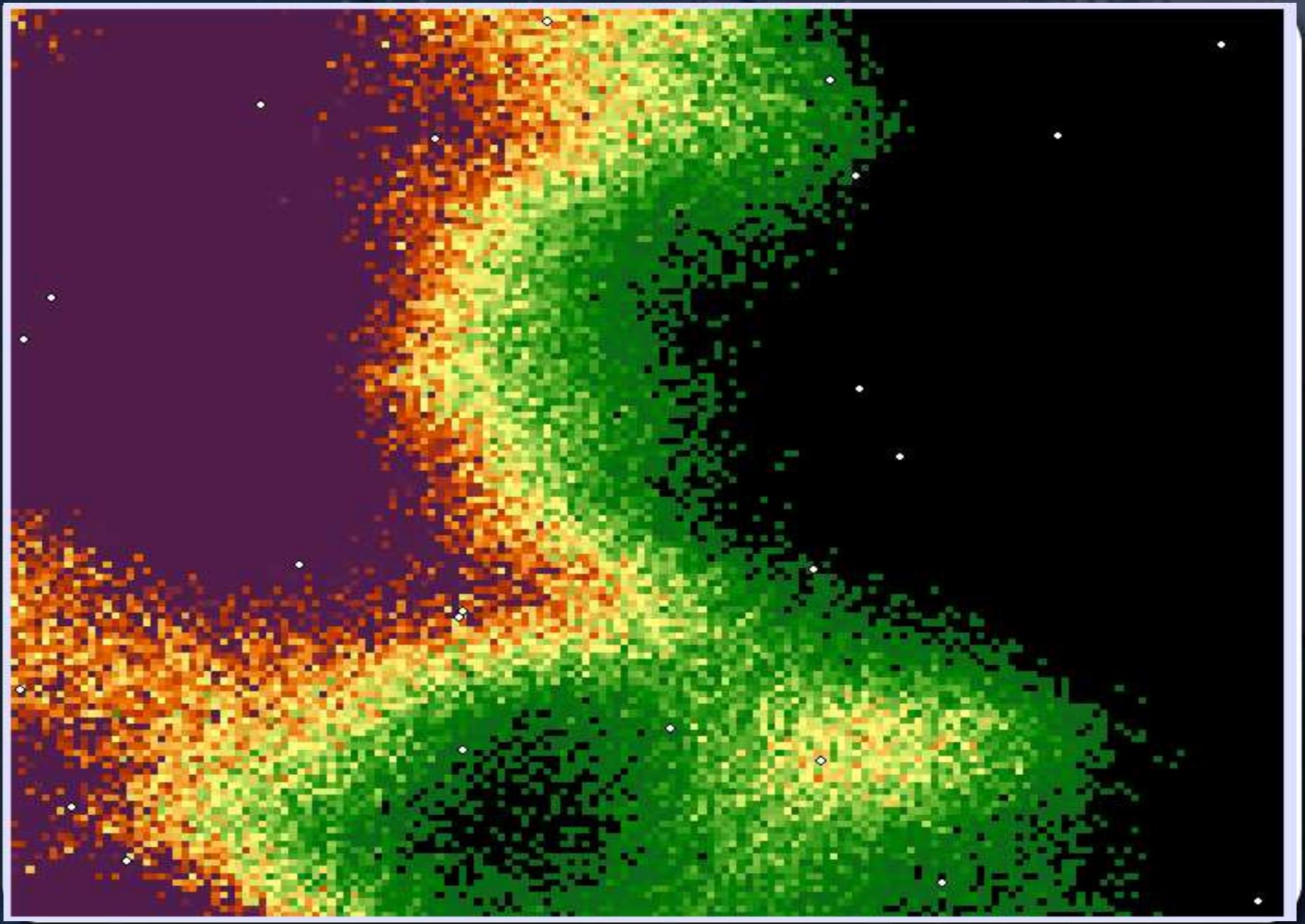


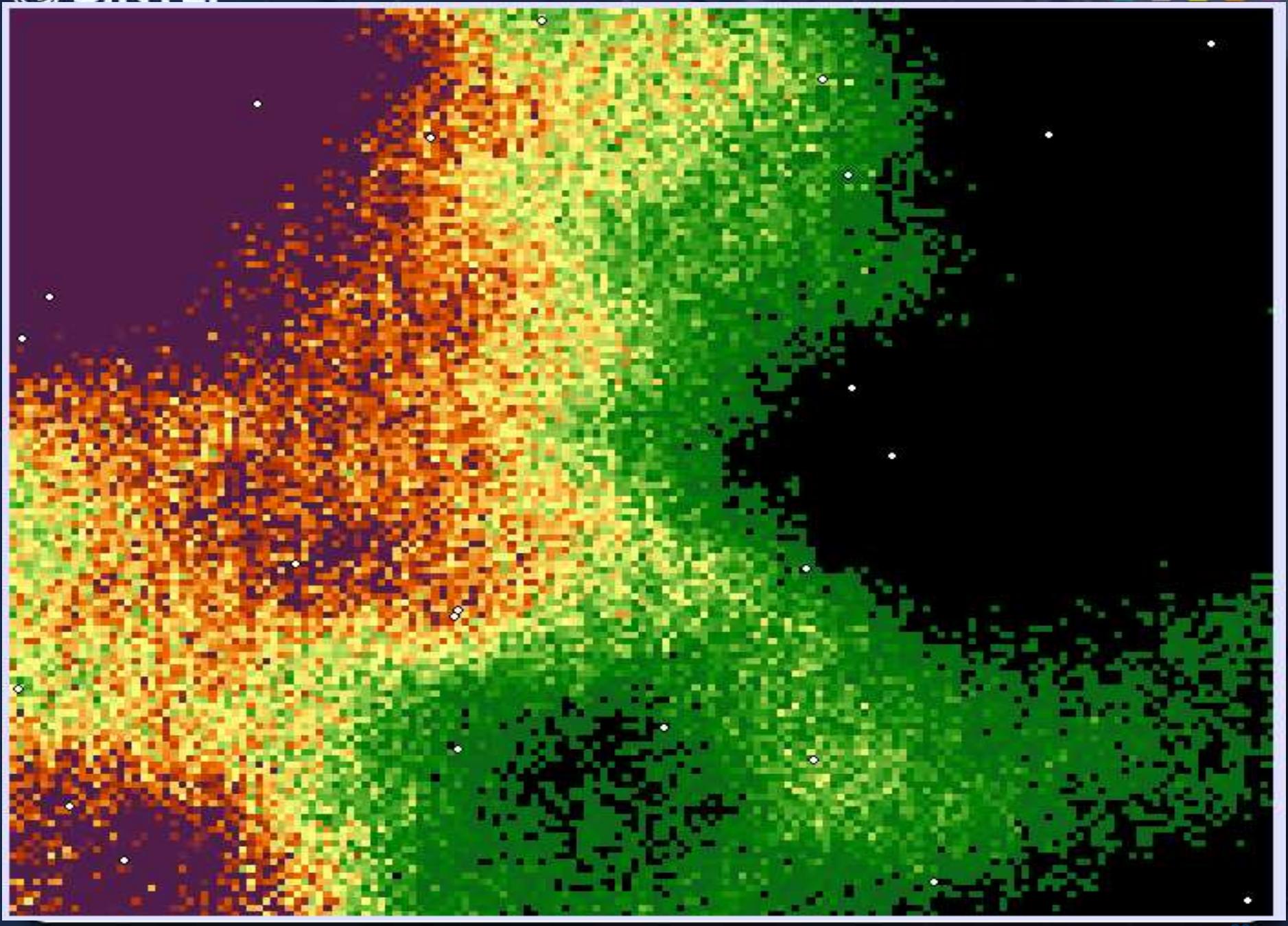


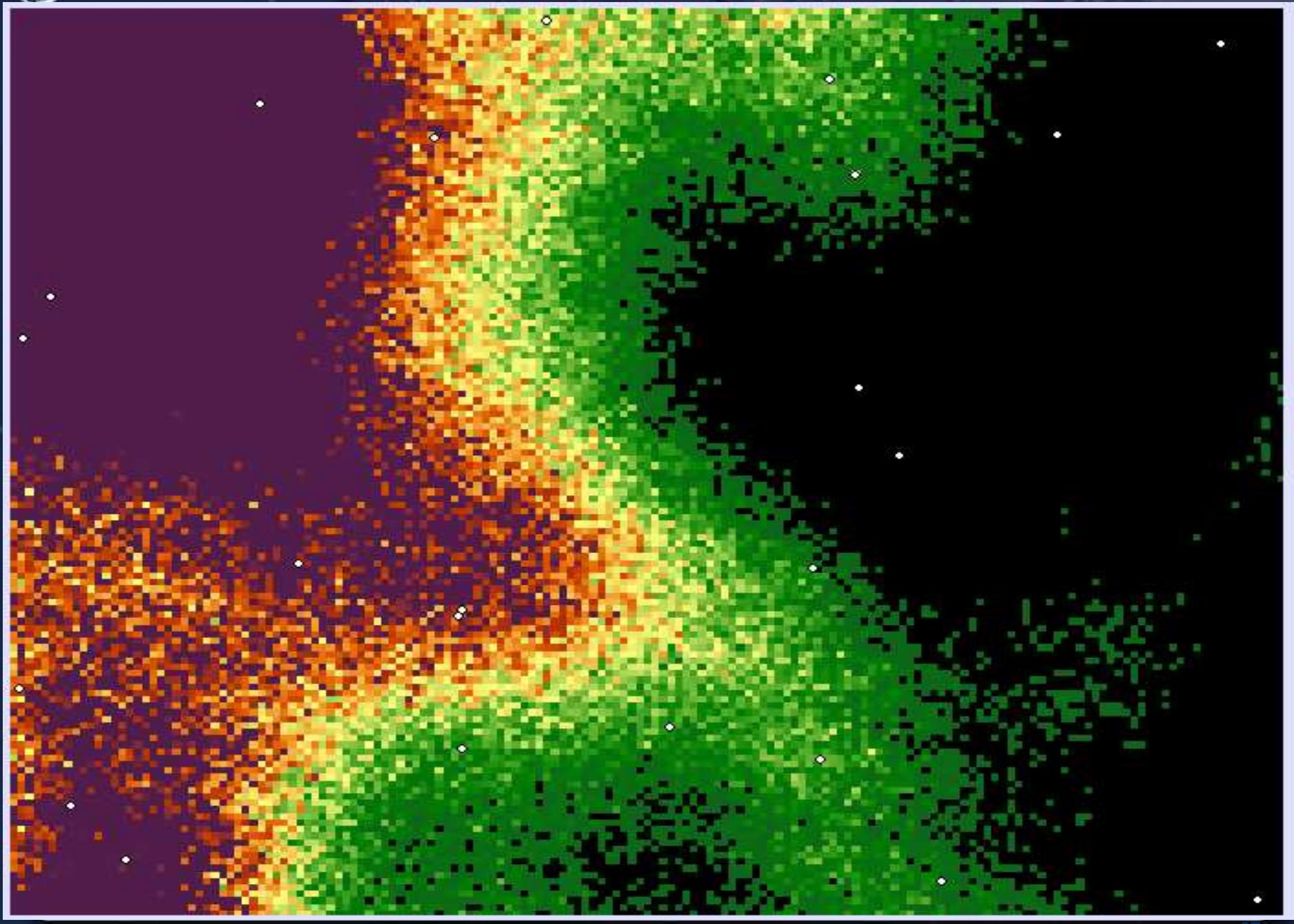


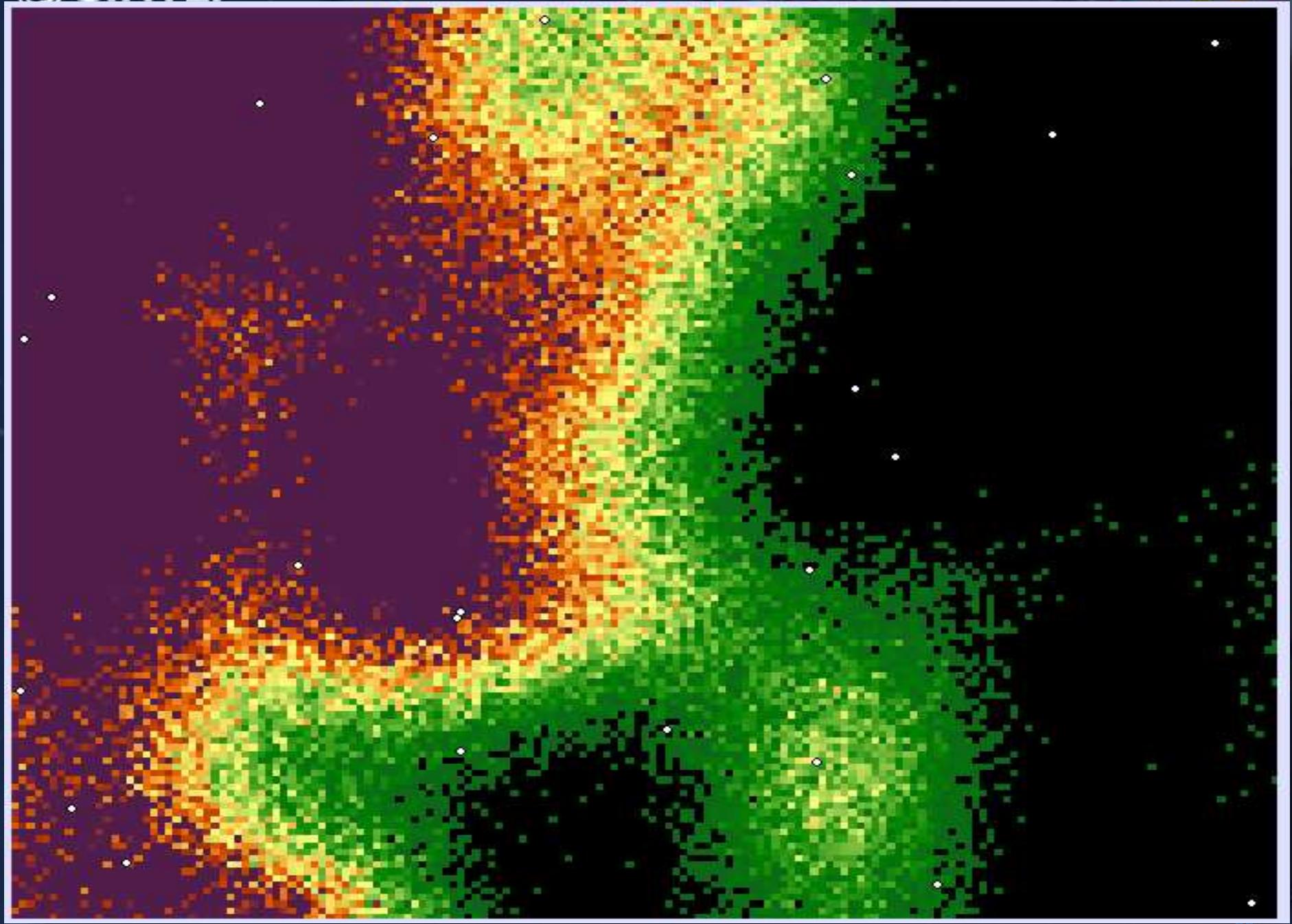


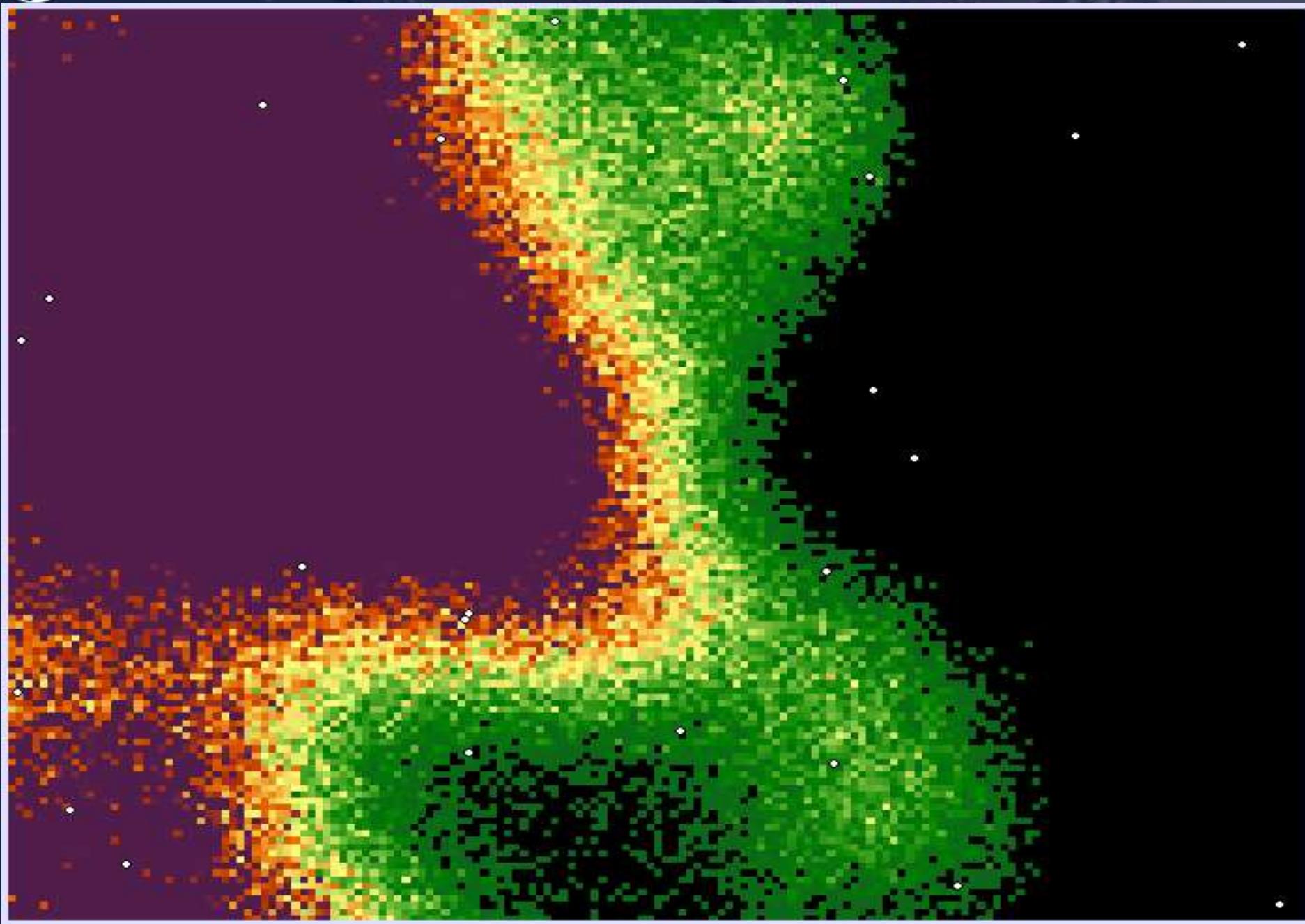


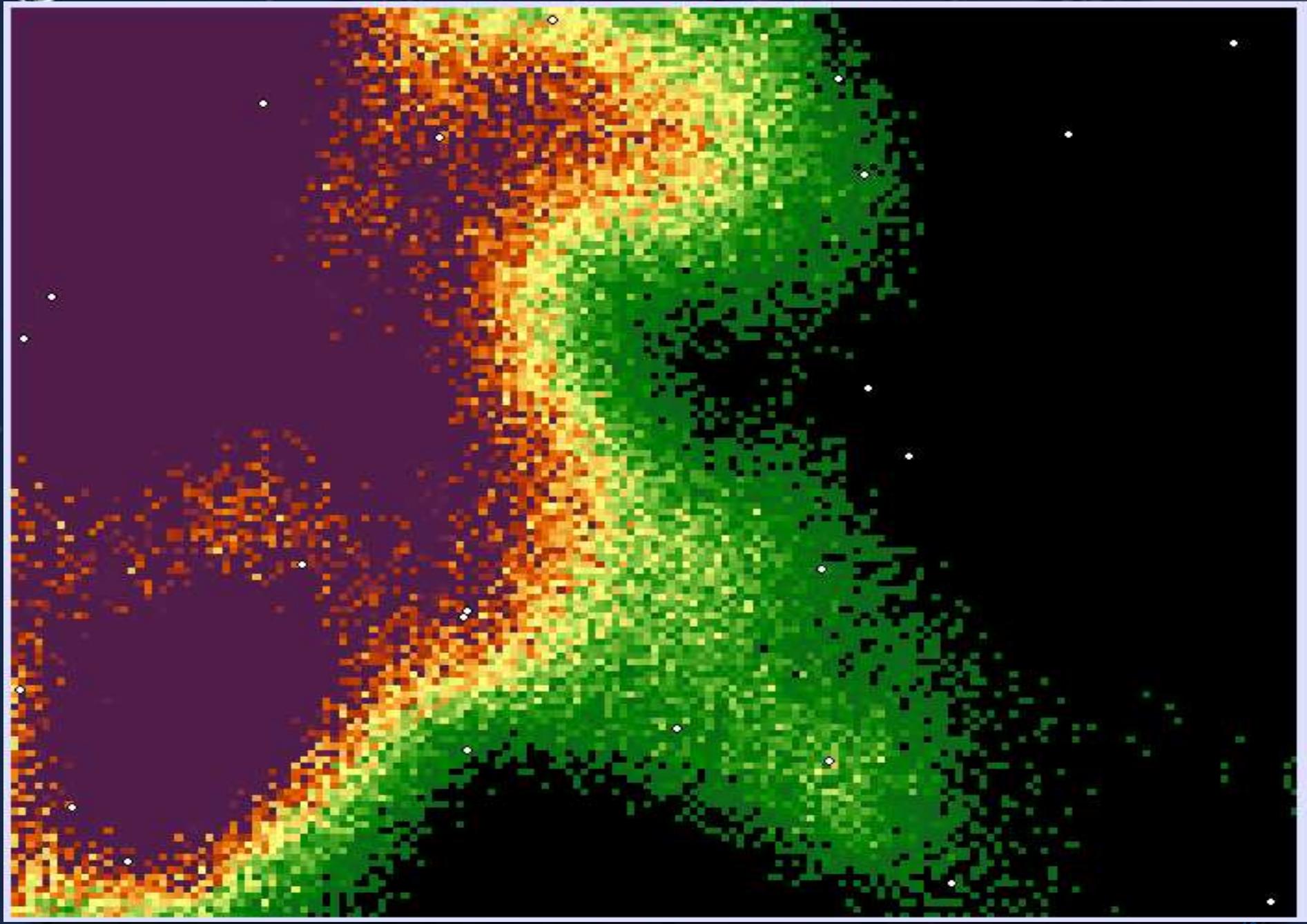


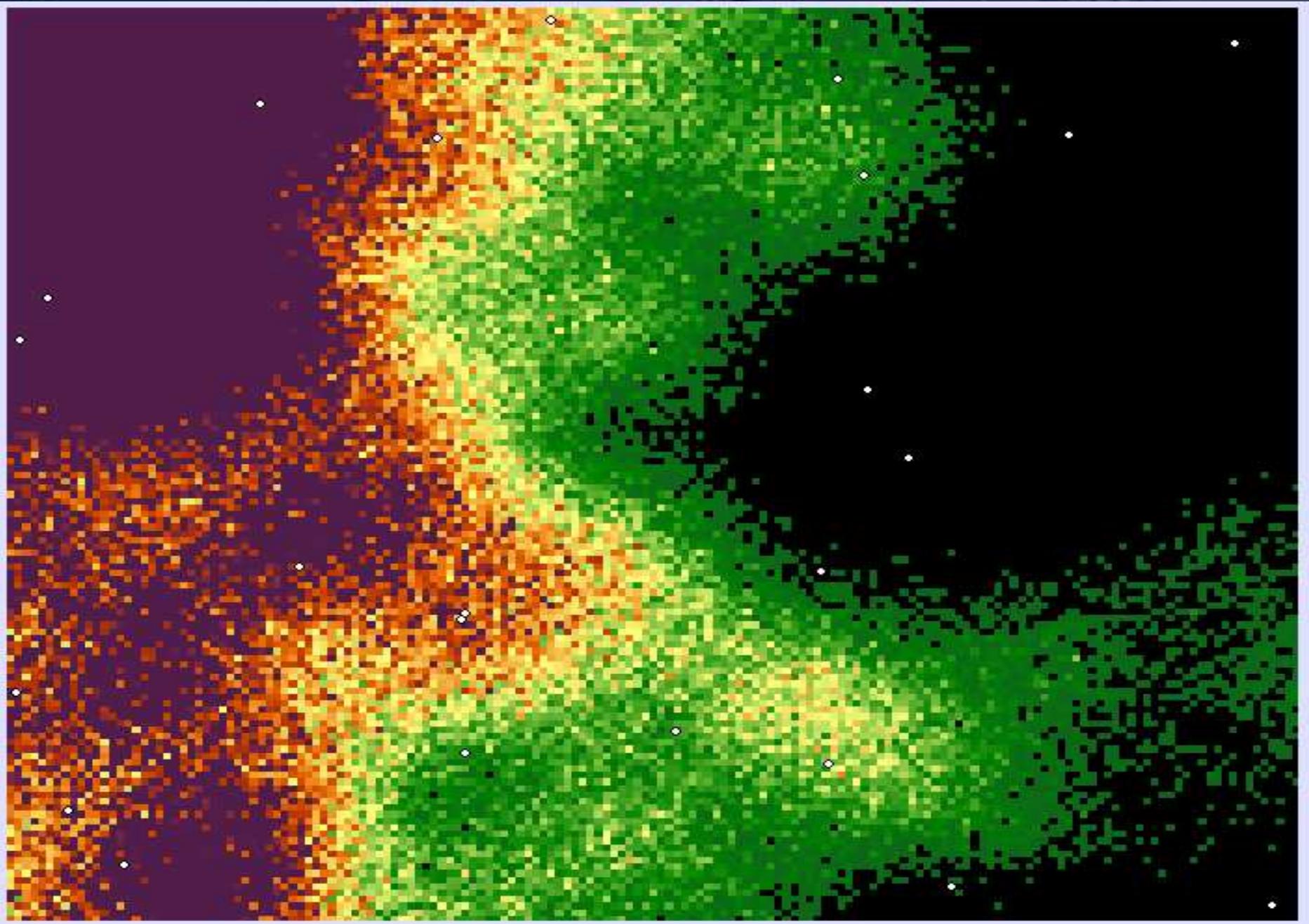


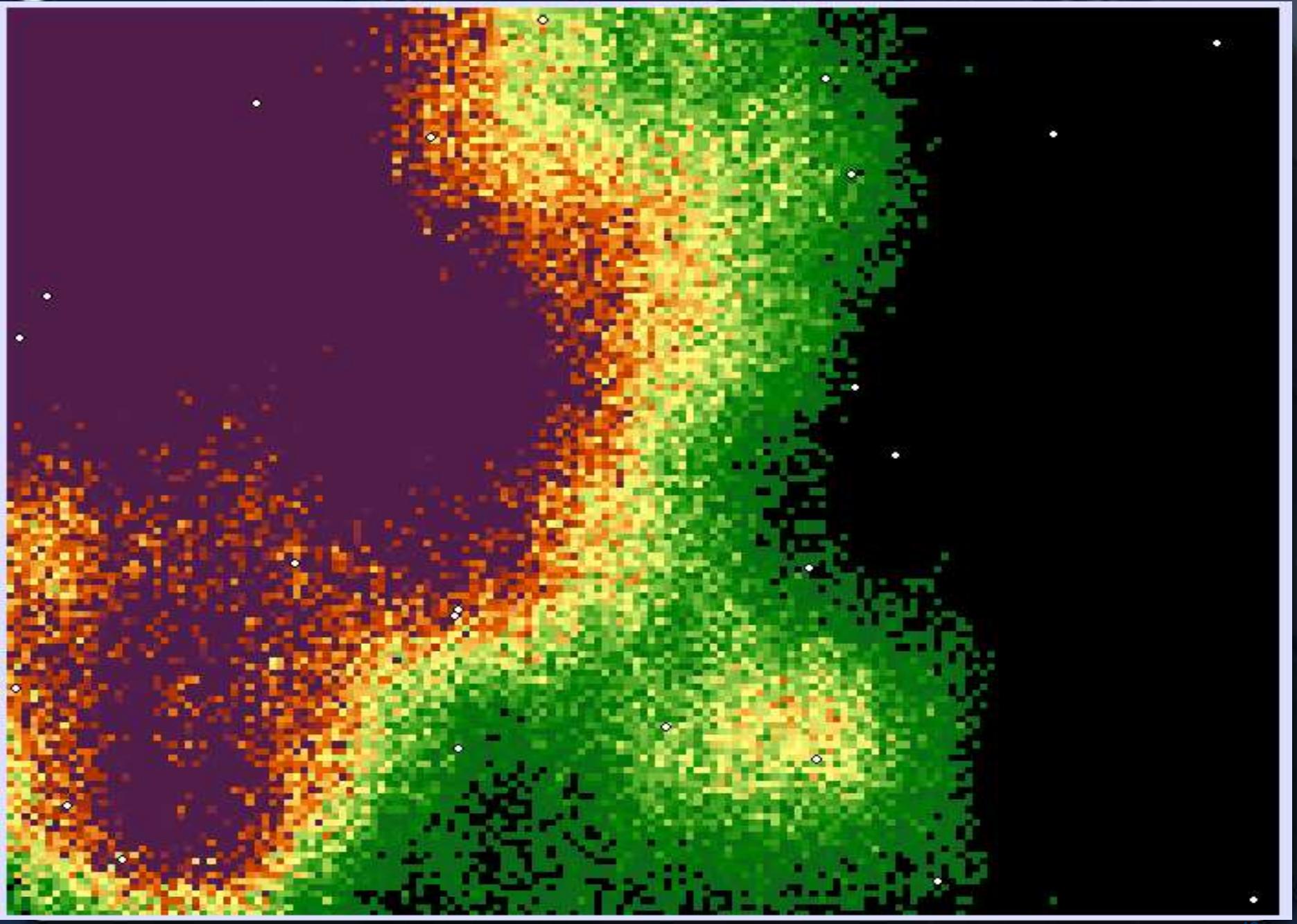


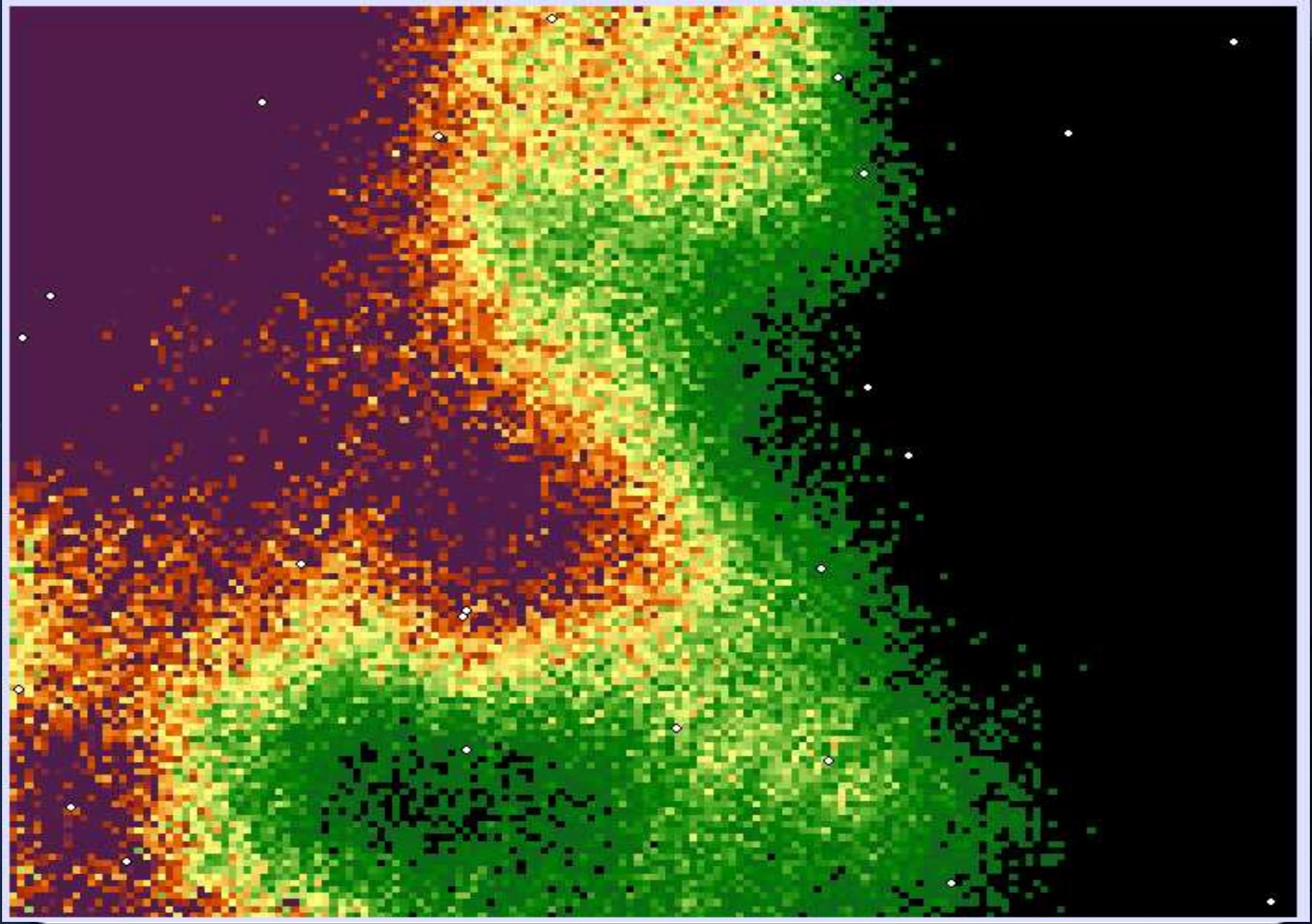


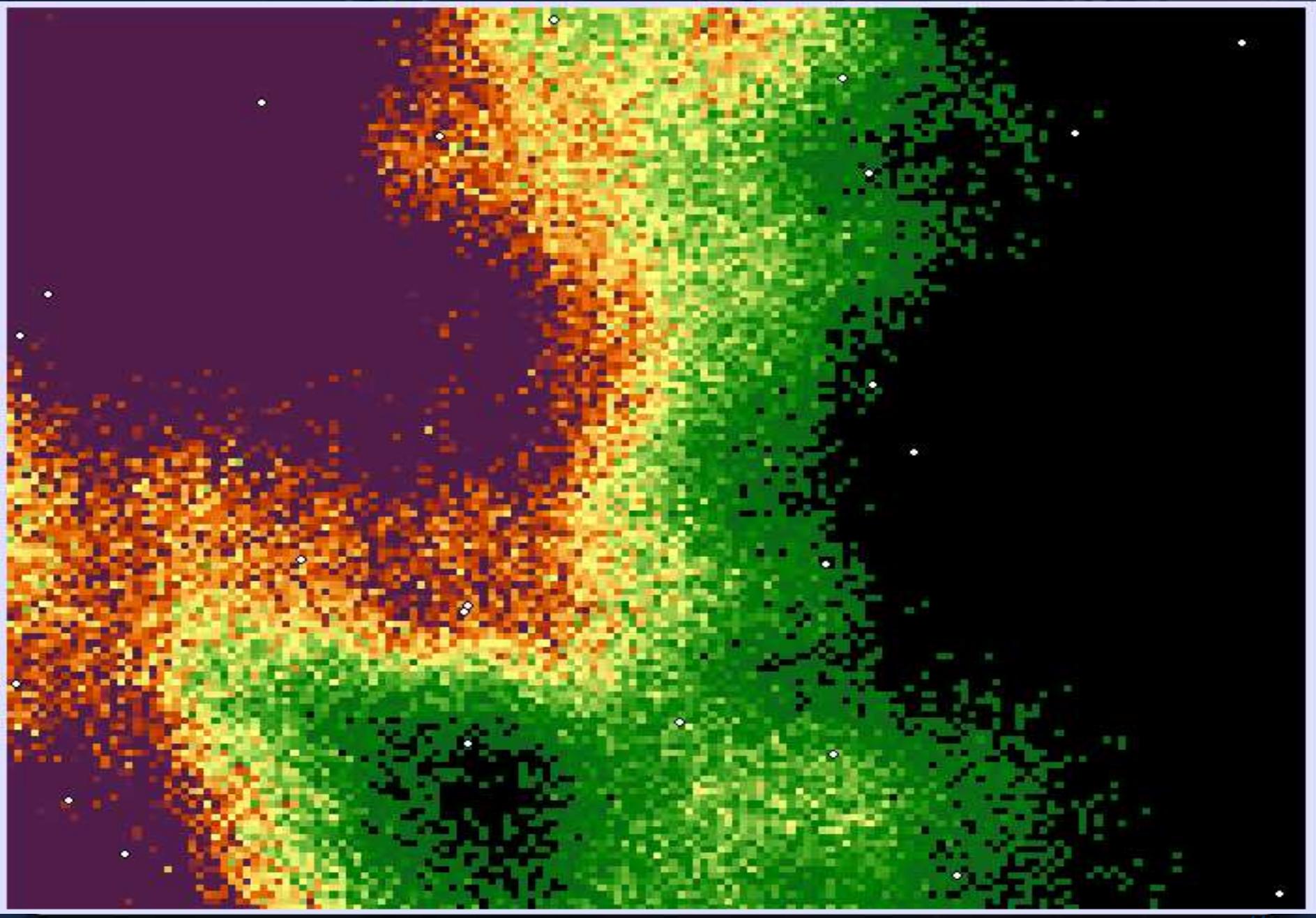


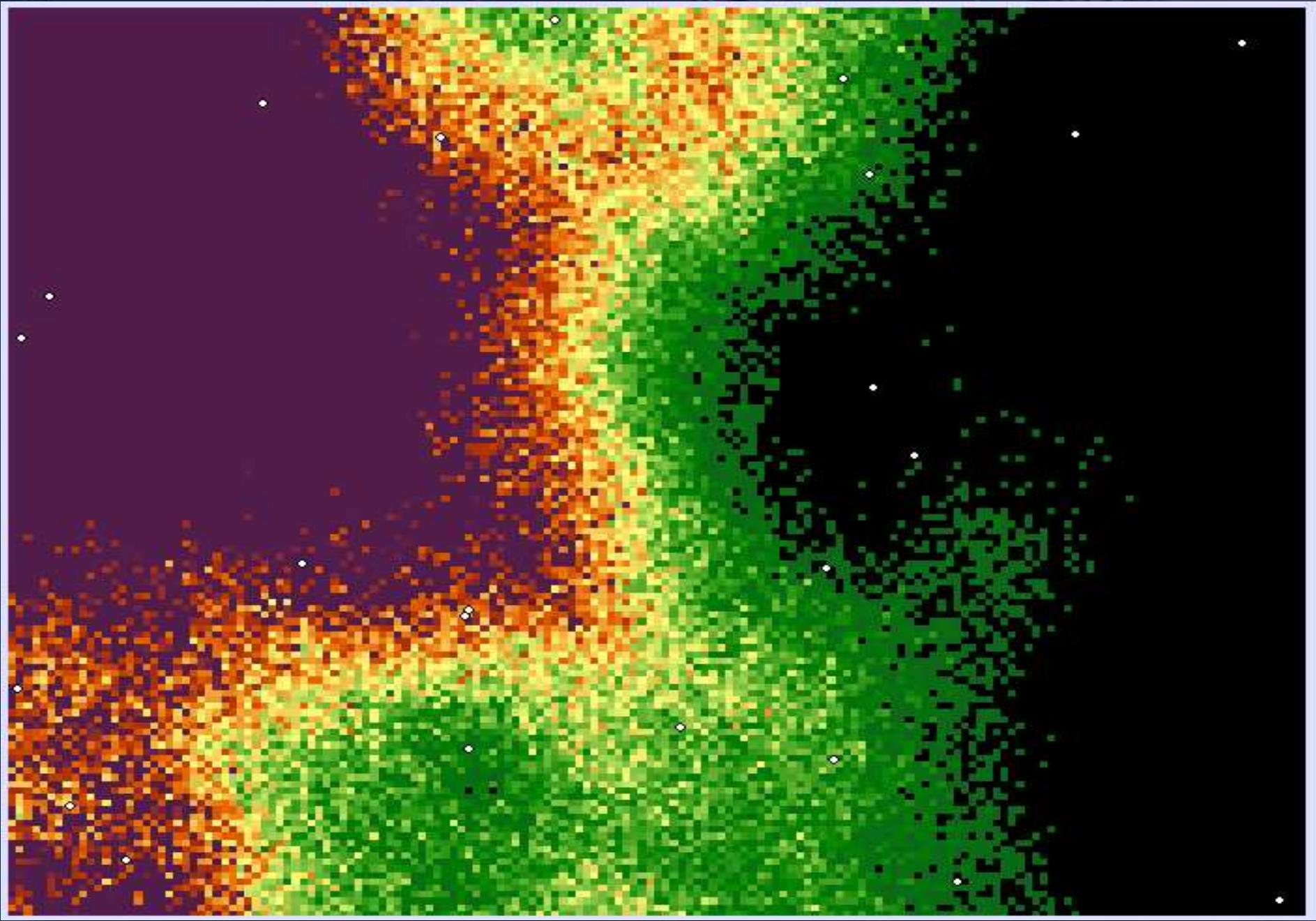


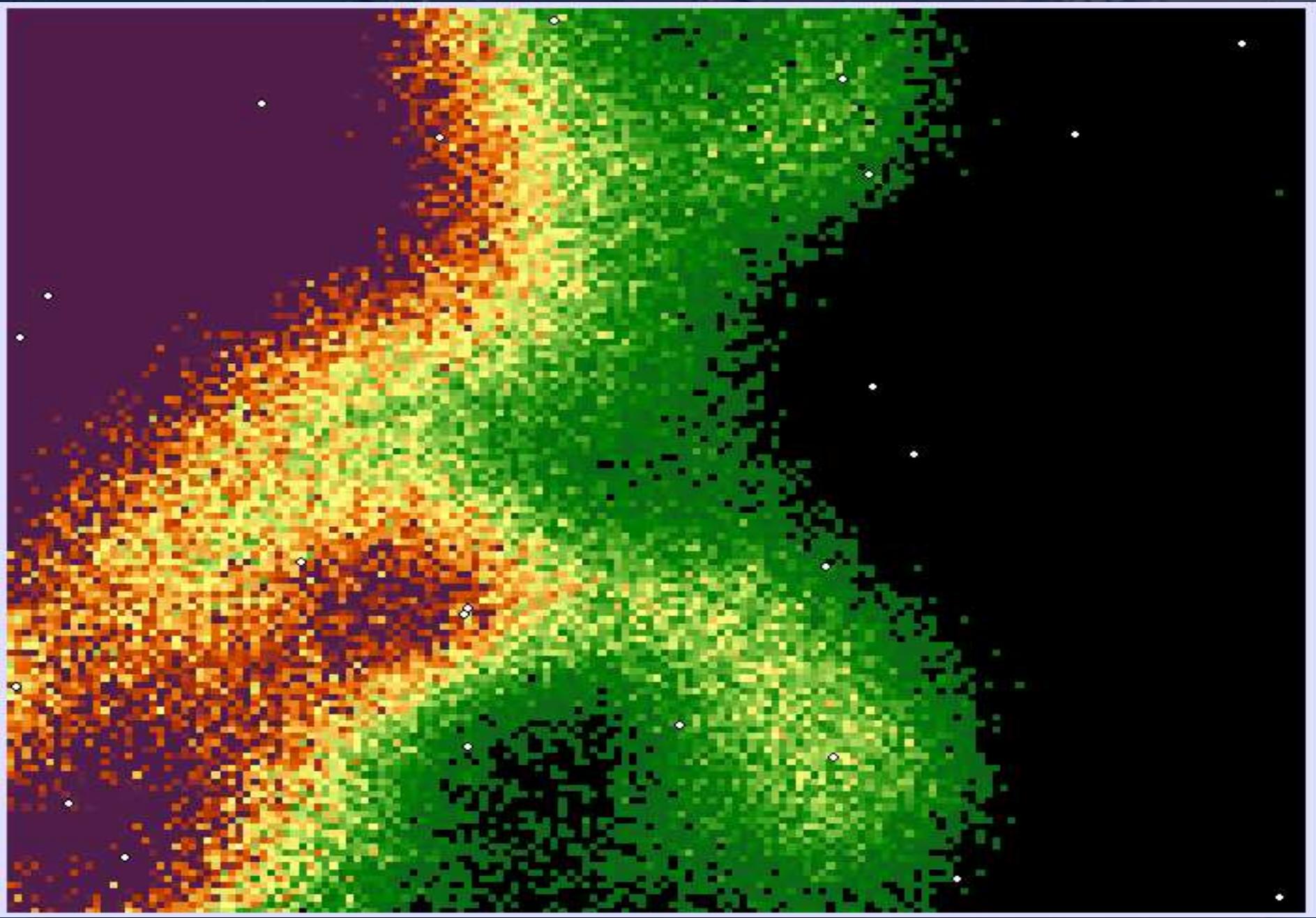




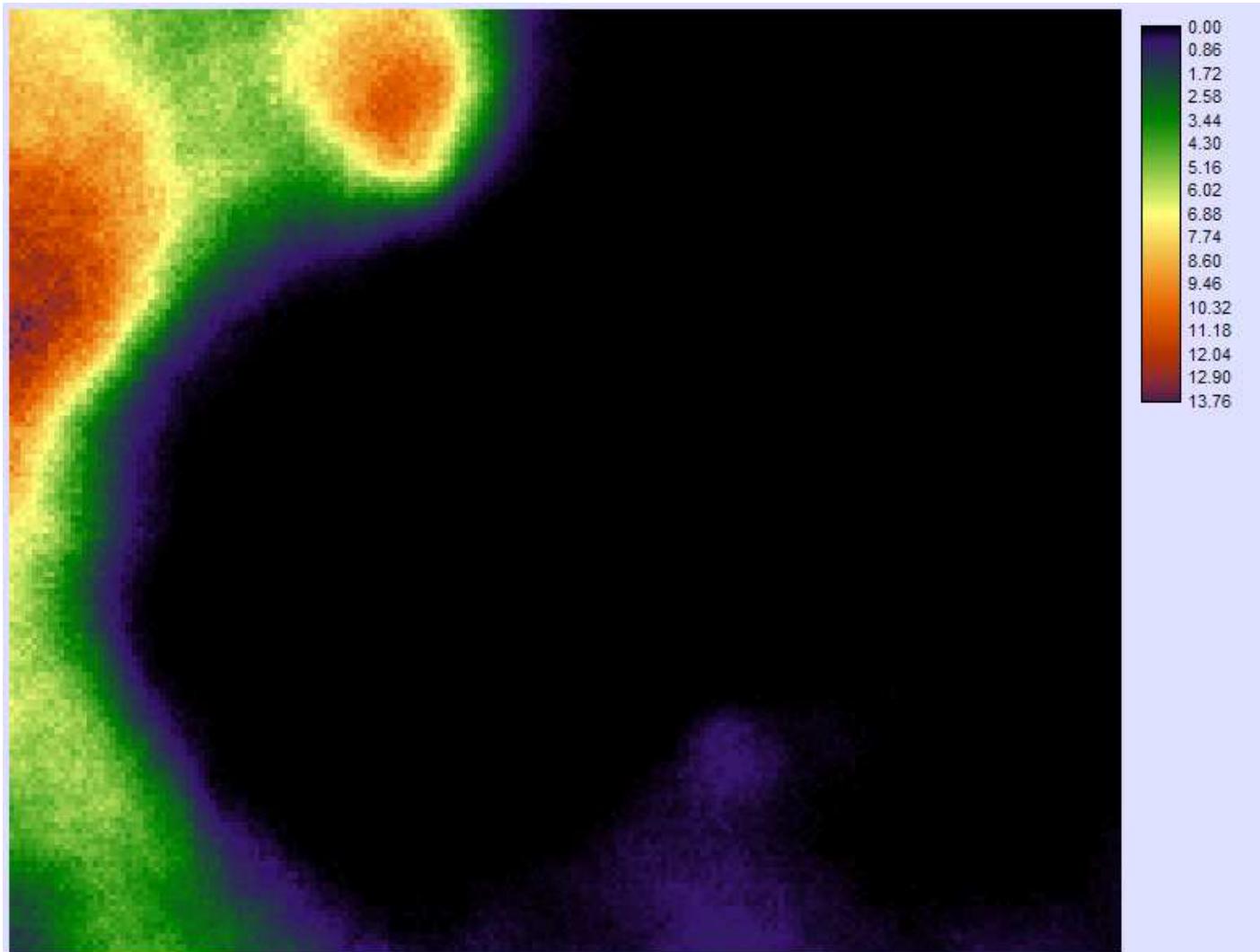




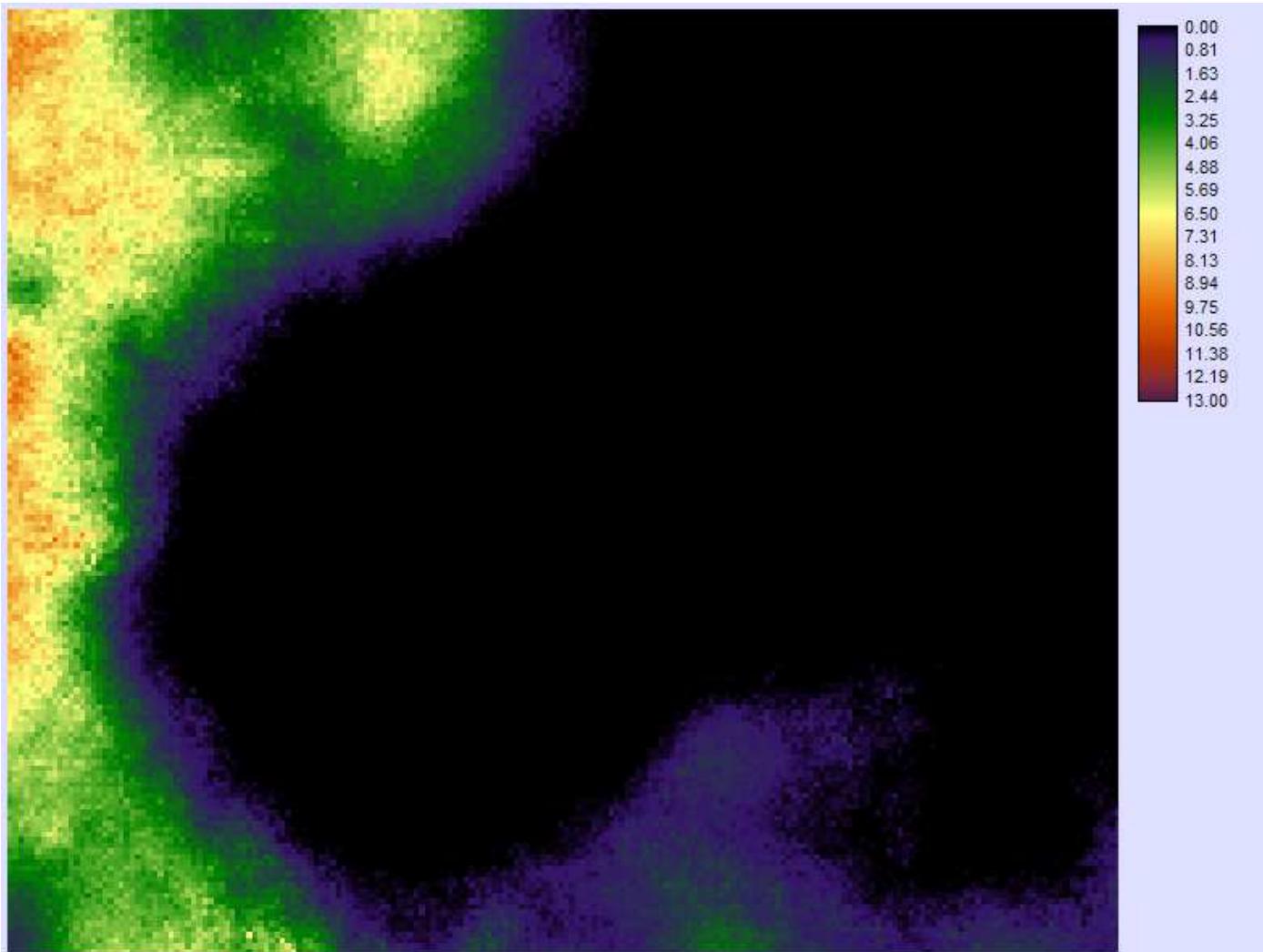




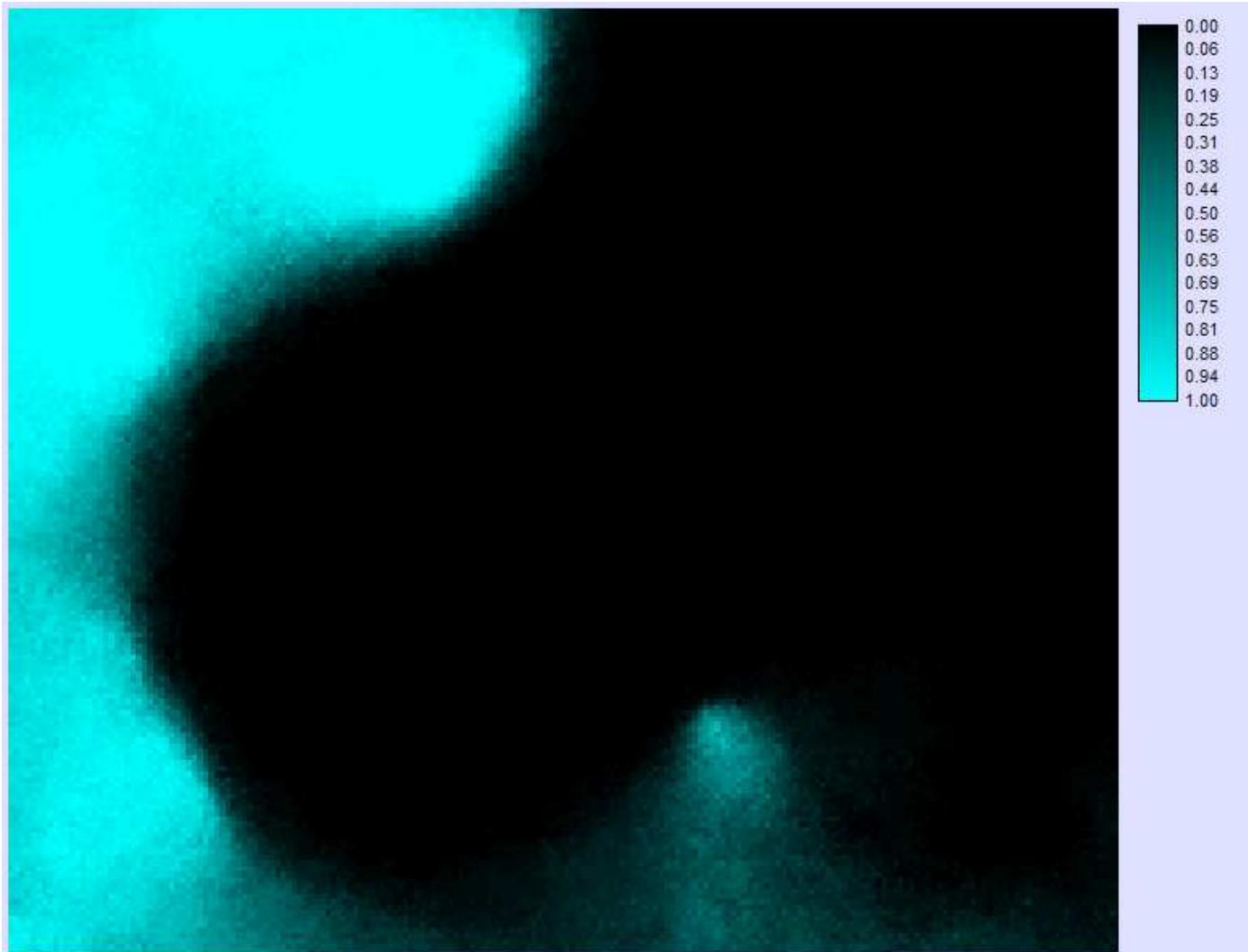
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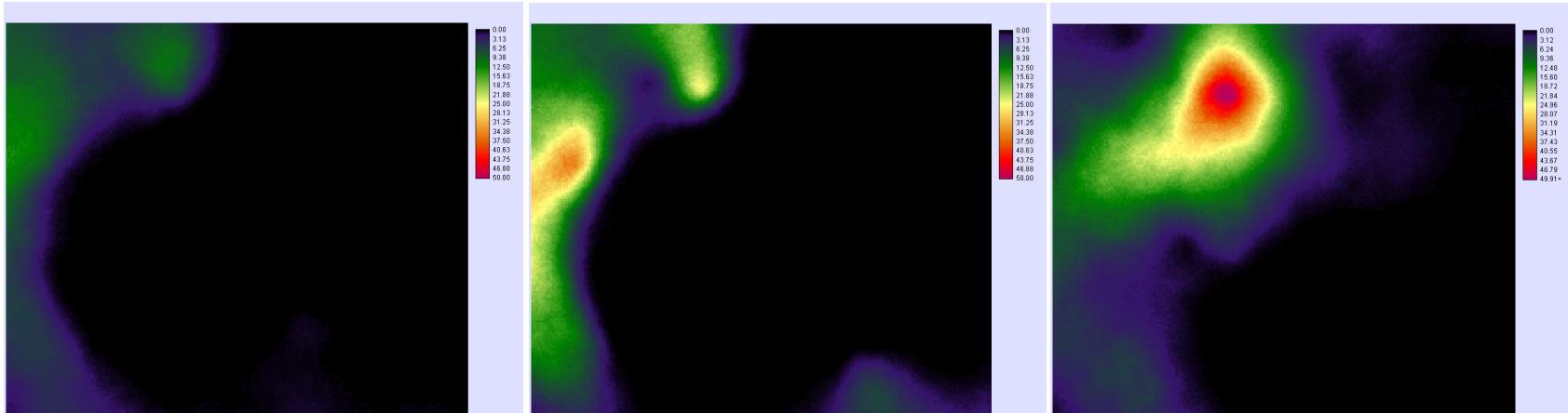
16. januar 2000, usikkerhet (SD):



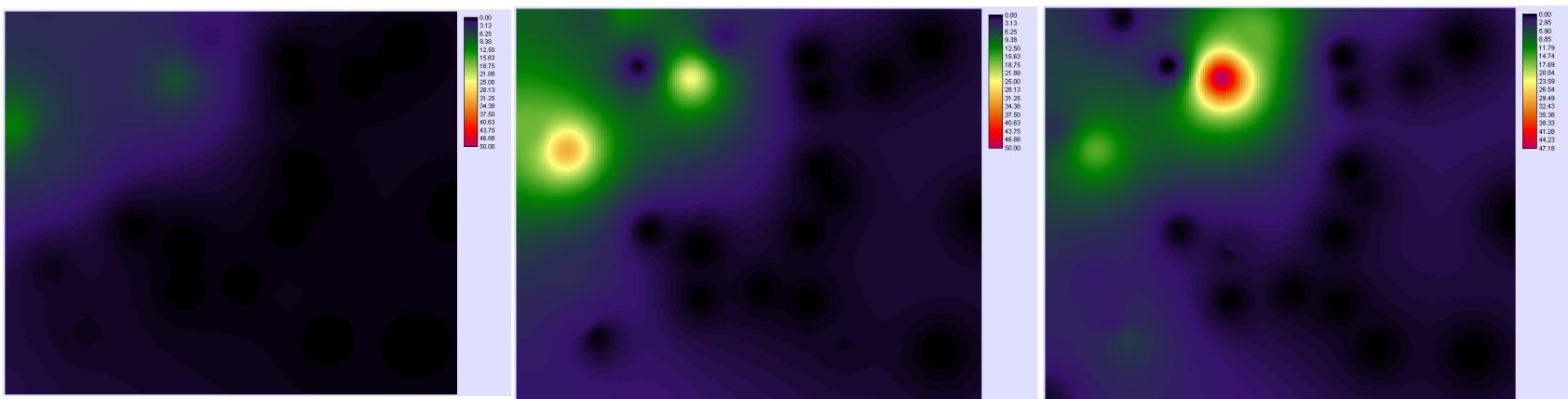
16. januar 2000, sannsynlighet for nedbør:



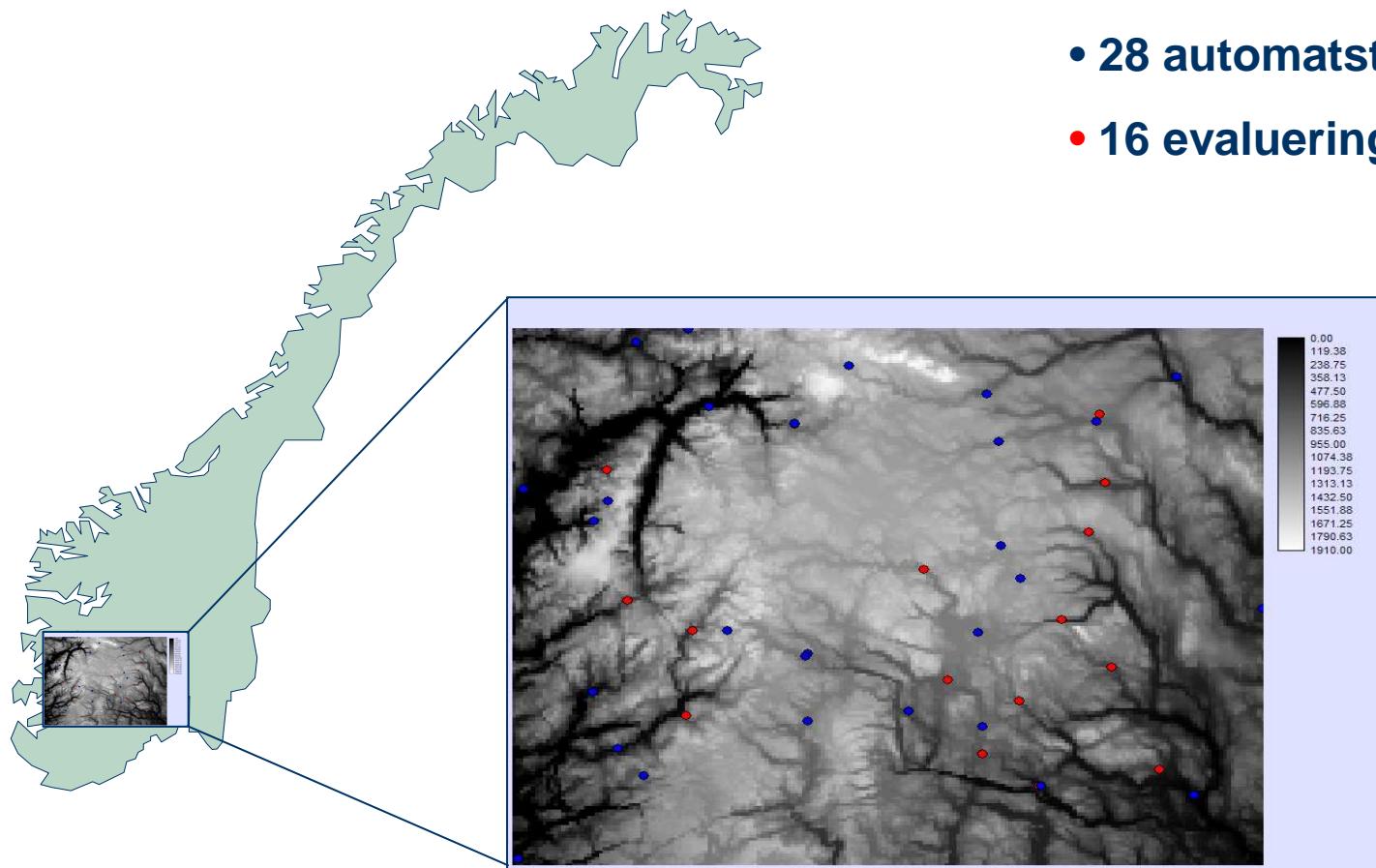
Simulation results (January 16 – 18 2000):



Inverse squared distance interpolation:



Eksempel: Hardangervidda



Evaluering:

